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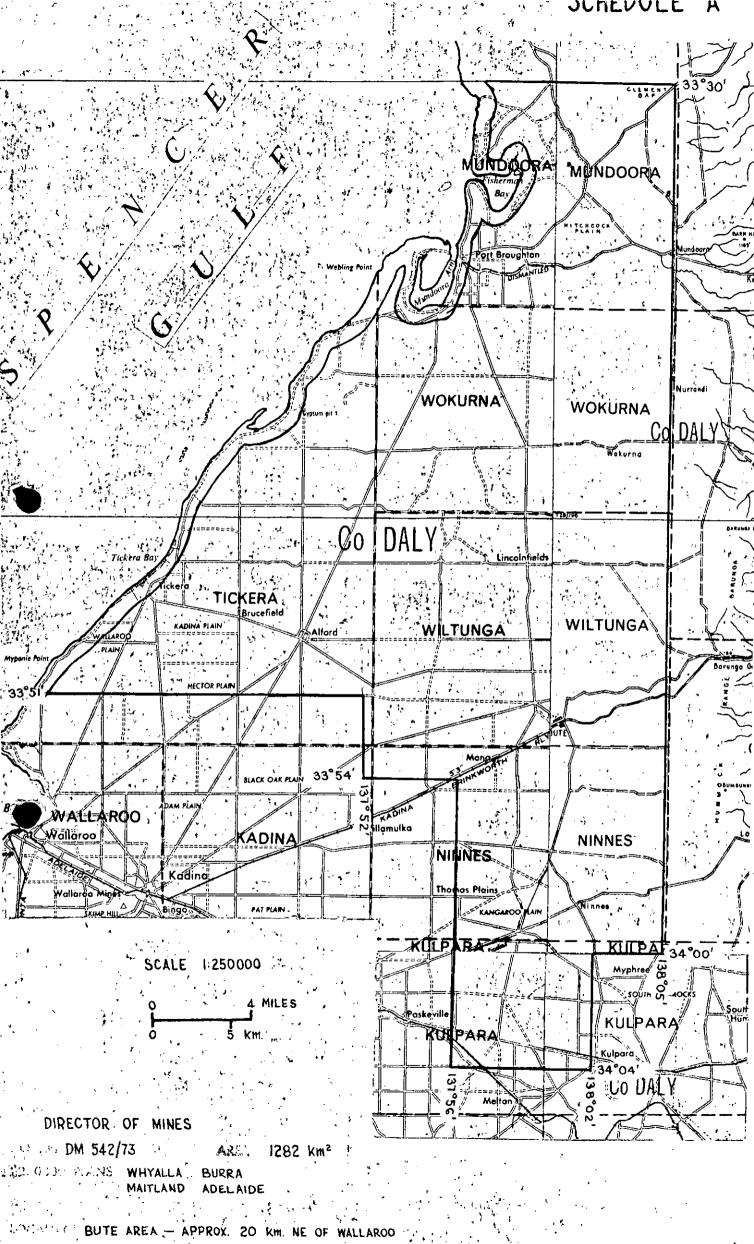
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# DEPARTMENT OF MINES SOUTH AUSTRALIA

## GEOTHYSICAL INVESTIGATION OF THE TICKERA ALECRD - FORT BROUGHTON AREA

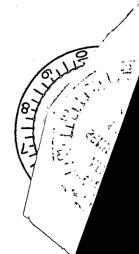
in Wallaroo and Broughton 1:63 360 sheet areas.

D.Y

# R.A. GIRDES Assistant Senior Geophysicist Exploration Geophysics Section

14th August, 1974

Rept.Bk.No. 74/463 G.S. No. 5482 D.M. No. 195/74



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# DEPARTMENT OF MINES SOUTH AUSTRALIA

Rept.Bk.No. 74/463 G.S. No. 5482 D.H. No. 195/74

#### TICKERA REGION

GEOPHYSICAL INVESTIGATION OF THE TICKERA-ALFORD-PORT BROUGHTON AREA in Vallaroo and Broughton 1:63 360 sheet areas

#### ABSTRACT

The Tickera-Alford-Port Broughton area located in Broughton and Wallaroo is an area of relatively shallow metamorphic rock, covered by a thin veneer of Quarternary and Tertiary sediments, which contain clays and highly saline layers. A systematic quantitative interpretation of Schlumberger Vertical Electrical Soundings was undertaken to delineate areas of high electrical transmissivity and the depth to the resistive basement as a feasibility study for future Electromagnetic and Induced Polarization surveys. A number of areas for future geophysical surveys have been outlined. Incoroporated in the report is an appraisal of previous geophysical data, namely Induced Polarization and Vertical Hagnetic Intensity, performed by Western Mining Corporation Ltd. in a relinquished portion of Special Mining. Lease No. 87.

#### 1. INTRODUCTION

The Tickera-Alford-Port Broughton ares shown in drawing No. S10352, is a region of shallow metamorphic or granitic basement covered by a thin veneer of Quaternary and Tertiary sediments. The main outcrops of metamorphic basement are restricted to the coastal section, south of Tickera township, see drawing No. 73-709.

The main geophysical problem in this region is caused by the high salinity horizons within the cover rocks, which have caused considerable difficulties for the

Induced Polarization curvoys conducted by Western Mining Corporation Ltd. and have caused 'skin effect' experienced by Sheard and Taylor (1973) in the area couth of the Fort Broughton Acromognetic Anomaly.

The Schlumberger, Vertical Blectrical Sounding Depth Probes (V.C.S.) arranged approximately on a 2000 yard grid (Willitery) were undertaken to establish the following:-

- (1) The depth and thickness of the low reststivity layers.
- (ii) The dopth to the top of the weathered metamorphic basement.
- (111) The depth to the high resistive basement, which correlates with either consolidated metamorphic basement or an intermediate consolidated horizon.
- (1v) To define the percentage of the area, which is not covered by low resistivity layers, thereby being able to establish the feasibility of conducting an airborne electromagnetic survey.

The V.J.J. survey was carried out between the 16th to the 19th April, 1973 and the 30th April to the 23rd May, 1973.

The interpreted results of the V.J.S. are compared later with the apparent resistivity results interpreted from the Induced Folarization resistivities for a 200 feet dipolodical configuration, recorded by Jestern Mining Corporation Ltd. (J.M.C.) in 1964/65.

The W.M.C. Induced Polarization frequency effects and vertical magnetic results were appraised to define possible likely target areas for mineralization investigations.

The area under investigation is shown in drawing

No. S10352 and the locations of the V.E.S. for both the

Tickera and Port Broughton surveys are shown in drawing No.

73-709, together with the W.M.C. Induced Polarization Coverage.

The vertical magnetic coverage and amgnetic contours, after

W.M.C. are shown in drawing No. 74-247.

#### 2. GEOLOGY

The geology of <u>Wallaroo</u> was compiled from photo interpretation by Crawford (1960) and included some detailed geological mapping of Jack (1917). The most recent compilation of the geology of <u>Broughton</u> is shown in the 1:250 000 preliminary sheet of WHYALLA.

W. McCallum (report in preparation) under the direction of B.P. Thomson remapped the <u>Broughton</u> and <u>Wallaroo</u> sheet areas within the current area of the Exploration Licence Area. The Concurrent subsurface geology after McCallum is shown in drawing No. 73-709.

## 5. PREVIOUS GEOPHYSICAL SURVEYS

# 3.1. REGIONAL DATA

#### A. Aeromagnetic

In 1952, the Bureau of Mineral Resources flew an aeromagnetic survey of Wallaroo at an elevation of 500 feet above ground level (a.g.l.) along lines orientated north-south and spaced 1 mile apart. This latter survey was flown as part of the exploration for iron in the northern portion

of Yorke Peninsula. Later, in 1960, the Bureau of Mineral Resources, whilst surveying BURRA and the eastern portion of UHYALLA, reflew Wallaroo, Wells (1962). This survey data was obtained along lines orientated east-west, with a flight line spacing of 1 mile, at an elevation of 500 feet a.g.l.

#### B. Regional Gravity

In 1967, the Exploration Geophysics Section of the Geological Survey of South Australia conducted a regional gravity survey of BURRA and the eastern portion of WHYALLA, with a station density of approximately 1 station per 16 square miles.

The Bouguer gravity contours in the Tickera area show a distinctive gravity low, probably explained by the presence of the Tickera granite, Jack (1917) intruding the Cleve Metamorphics (?).

The station density of this regional data is unsatisfactory for the exploration and delineation of geophysical structures within an area of shallow metamorphic terrains.

## 3.2. DETAILED GEOPHYSICAL DATA

## A. <u>Vertical Magnetics and Induced Polarization</u>

During the period 1963-1965, Western Mining Corporation Ltd. (W.M.C.) undertook reconnaissance vertical magnetic and Induced Polarization surveys of the whole of Special Mining Lease 87. These reconnaissance surveys consisted of a number of lines orientated north-south, approximately % mile apart. The geophysical coverage of this geophysical data

is shown in drawing Nos. 74-247 and 73-709, for the Magnetic and Induced Polarization data respectively. The vertical magnetic intensity contours after W.M.C. are shown in drawing No. 74-247.

Details of the geophysical data of this area is tabulated below for both the Magnetics and Induced Polarization.

•	MAGNETICS	INDUCED POLARIZATION
Number of Lines	38.	16
Station Interval	300 feet	200 feet (dipole-dipole)
Total Mileage Covered	221.72 miles	72.23 miles
Sea of the	(384.1 km)	(125.13 km)

Approximate Area Covered 78.0 square miles.

A reinterpretation of this data is given later in this report.

#### B. Electrical Vertical Soundings

In December 1972, the Exploration Geophysics Section of the South Australian Geological Survey undertook a resistivity survey of the Port Broughton Aeromagnetic Anomaly, Sheard (1973) unpublished report. This survey consisted of 24 V.E.S. of which 50% are unsatisfactory for depth interpretation due to:

- (a) The skin effect produced by highly conductive overburden for a frequency of 3 Hertz; and
- (b) the comparatively small current electrode half spacings AB/2 distance used.

The remaining V.E.S. using frequencies of 0.3 and 0.1 Hertz, have been reinterpreted and are incorporated in this report.

#### 4. REINTERPRETATION OF PREVIOUS GEOPHYSICAL DATA

The reinterpretation of the previous geophysical data is restricted to the reconnaissance Induced Polarization and Vertical Magnetic Traverses surveyed by Western Mining Corporation Ltd. in the Tickera area, located immediately north of Exploration Licence No. 32 held by North Broken Hill Pty. Ltd.

#### 4-1.1. Induced Polarization Traverse Data

The I.P. data available from the V.M.C. reconnaissance traverses is in the form of diagonal plotted pseudo-sections of the apparent resistivity, metal factors, and uncontoured frequency effect data, for a dipole-dipole configuration.

The dipole spacing 'a' used was 200 feet, for n=1, and 2 and 3, where 'na' is the smallest distance between the transmitting and receiving dipoles.

Line profiles of apparent resisitivity and percentage frequency effect were replotted with respectively logarithmic and linear scales for the vertical axis for the condition where the number of dipole separations (n) - 3 for each traverse.

A value of n = 3, was chosen to minimize local variations of the upper cover rocks presenting an apparent depth of penetration of approximately 400 feet (122 m) for the pseudosections. The apparent resistivity at this depth should reflect:

- (a) the regional scale undulations of the bedrock surface probably filled with saline material, and
- (b) lithological changes within the basement.

The frequency results should reflect broad lithological units and possible local frequency effects worthy of reinvestigation.

The interpretation of the profile data of the apparent resistivity and frequency effect will be discussed below.

#### 4.1-2 Apparent Resistivity Data

The inspection of the apparent resistivity logarithmiclinear plots of each traverse showed a considerable amount of high frequency noise, due to statistically distributed reading errors or surface geological noise. The latter is produced by either the positions of the dipole electrodes with respect to lateral changes in resistivity of near surface bedrock, or, rapid changes in depth of overburden, as shown by Van Nostrend et. al., (1966)

This profile data was smoothed and contour intervals of 10, 20, 50, 75 and 100 ohm metres were selected for the convenience of compiling a contour map of the area, shown in drawing No. 73-710.

The areas showing resistivities of less than 20 ohm metres are considered to be probably areas of deep overburden containing saline layers. The areas with apparent resistivity of 20 to 50 ohm metres are considered to represent areas with either thin saline layers or low resistive basement. The areas indicated as greater than 50 ohm metres probably correlate with either basement highs or areas of high resistive basement.

#### 4.1-3 Percentage Frequency Effect Data

The inspection of the percentage frequency effect of each traverse showed a considerable amount of statistical and geological noise. The major frequency effects were located at the northern ends of the eastern lines and occurred probably due to electromagnetic coupling. Other frequency effects were also present, and it was considered necessary to estimate their degree of significance with respect to the surrounding noise level, and to relate these anomaly values with the magnetic and resistivity data.

Frequency effects were first classified on whether they were statistically significant considering the magnitude of local statistical noise envelopes. This principle has been applied to outboard radiometrics in Western Australia, Tipper and Gerdes (1971).

Significant I.P. anomalies shown in drawing No. 73-708 were classified into four types, based on their correlation with the magnetics and apparent resistivities, shown below:

#### INDUCED POLARIZATION ANOMALY TYPE

# TYPE COLLENTS AND CLASSIFICATION 1.P. anomaly associated with a magnetic peak.

- B I.P. anomaly located on the flank of a magnetic anomaly.
- C I.P. anomaly associated with a low resistivity anomaly.
- D I.P. anomaly possibly associated with man-made feature.

The type A, I.P. anomaly associated with a magnetic feature has been known to correlate with graphite and magnetite in the area south of this region.

The type B, I.P. anomaly may correlate to graphite, but has been known to correlate with amphibolites in this region.

The type C, I.P. anomaly may be explained by saline layers but should not be discounted as a possible zone of mineralization.

The type D, I.P. anomaly has been found in this area to relate to water mains and pipes, and high tension power lines, the latter being located just south of Alford. It is interesting to note, that the I.P. effect of the water pipe along the southern boundary of the survey area is intermittent, and similarly those located on the Tickera-Wallaroo road. The larger frequency effects should be resurveyed to verify the existence and non existence of I.P. effects associated with these water pipes.

The frequency anomalies of importance are generally of type B and perhaps of type A.

## 4.2. VERTICAL MAGNETIC INTENSITY TRAVERSE DATA

The profiles of the vertical magnetic intensity traverse data were plotted on an extended sheet, shown in drawing No. 74-107, for better resolution. Original data contoured at an interval of 250 grammas, by W.M.C. is shown in drawing No. 74-247, and indicates some false trend directions.

An inspection of these profiles showed that it was very difficult to establish definite trends, as this data was considerable noisy.

The method adopted was to subdivide the data into magnetic zone types shown in drawing No. 74-454, and these zone types would define the overall grain of the basement rocks.

#### The Magnetic Zones and their significance

The magnetic zone boundaries were located at the position of the outermost inflexion point of a set of magnetic anomalies with most, but not necessarily all, of the anomalies in any zone being of that type. The zone type classification and anomaly characterisitic are given in table 1.

#### TABLE 1

Zone type	one type Amplitude Range		Characteristic	
1	less than 50 gammas	Poor	Linearity	
2	50 to 200 gammas	ti .	Ħ	
3	200 to 300 gammas	Ħ	<b>11</b>	
4	300 to 500 gammas	iz	• <b>11</b>	
5	500 to 750 gammas	, 11	. 11	
6	750 to 1000 gammas	n	W.	
7	1000 to 2000 gammas	11	ń	
.8	greater than 2000 gammas	n	n	

Type 1-zones are interpreted as being either non-ferruginous metasedimentary rocks or near homogenous acidic igneous and/or gneissic masses. Included in this zone type are areas of possible relatively deep basins i.e. up to 300 m deep filled with Tertiary and Quaternary sediments.

- Type 2 zones are interpreted as being slightly more magnetic than type 1 zones and correlate with the so called 'Tickera granite'.
- Type 3 zones are considered as being areas of weakly magnetic sediments and 'intermediate' type gneisses
  and schists. The zone type 4 is slightly more
  magnetic than type 3 zones.
- Types 5 and 6 zones are considered as being areas of magnetic metasediments, and meta-basic rocks.
- Types 7 and 8 zones are interpreted to correlate with iron rich metasediments and/or iron rock meta-basics.

The zone types 6 and 7 in the southwestern corner are considered to correlate with hornblende schists and hornblende graphite schist. The zone type 8 located in the north of the area correlates with a magnetite - quartz rich rock, assumed rock source of the Port Broughton Aeromagnetic Anomaly.

#### 5. SURVEY PROCEDURES FOR THE VERTICAL ELECTRICAL SOUNDINGS

The electrical resistivity method adopted for this investigation was the Schlumberger electrode configuration for the vertical electrical sounding (V.E.S.). This system consists of four electrode positions placed in a line and located symmetrically about a centre point. The current, generated by the Landrover, and modified by the Geoscience Induced Polarization, Lower Power Transmitter, operating at 3, 0.3, 0.1 Hertz, was passed into the ground by means of the outer electrodes A and B. The 0.3 and 0.1 Hertz were used to evercome the 'skin effect' caused by a low resistivity layer.

The resulting voltage was measured by a McPhar P660 Induced Polarization Receiver between the two inner potential electrodes M and N.

The field data for the Schlumberger V.E.S. had
the electrodes spaced so that the distance between the two
centre potential electrodes MN/2 is less than 20% the distance
the outer current electrode pair AB/2. For the V.E.S. the
outer current electrodes are progressively moved farther
away from the centre, whilst the inner potential electrodes
are moved only when the voltage between them becomes small.
The electrode half spacings for the potential electrodes MN/2
and current electrodes AB/2 are given in Gerdes et.al., (1973).

#### 6. REDUCTION AND METHOD OF INTERPRETATION OF THE V.E.S. DATA

The apparent resistivity for each particular electrode spacing was calculated, based on the formula for the Schlumberger configuration, shown in drawing No. S 9364. The apparent resistivity in ohm metres was plotted on bilogarithmic graph paper as a function of half the current electrode spacing AB/2 (in metres). Each curve was then interpreted by a curve matching of the Standard 2 and 3 layer master curves prepared by Orellana and Mooney (1966), using the Auxillary Point Method, advocated by Zohdy (1965). The latter method permits a substantial reduction in the required number of master curves.

The interpreted model, assuming horizontal homogeneous and isotropic layers consisting of layer thickness and resistivities, were subjected to a computer programme called AUTORESCUR developed by G. Pilkington (Mathematical Geophysicist) in Fortran 4. This produced better thickness and resistivity

values for the particular layers of the specified model.

A better solution was obtained for approximately 60% of
the V.E.S. It was found for the remaining 40% the major
draw back of AUTORESCUR was that for steeply descending
curves interpreted as double descending Q types curves, the
computer results were unable to improve the solution. In
this case, the computer solution was unsatisfactory, and the
original model was in most cases a better fit, when the
RESCUR plots were compared with the field curve.

In general, the AUTORESCUR programme is unable to accommodate resistivity contrasts greater than 200:1.

The models of the AUTORESCUR results and some original models were then processed using another computer programme called RESCUR, also developed by the above, in Fortran 4; see Pilkington (1973) for details of programme. The results from this programme enabled a direct comparison of the interpreted theoretical model curve to be compared with the field curve.

# 7. INTERPRETATION OF THE V.E.S. RESULTS

The interpretation of the V.E.S. assumes that the layers can be represented by a series of homogeneous horizontal layers. Small dip components up to 15 degrees, can be considered to approximate the horizontal layers if the subsurface of the dipping bed is planar, for the case of a schlumberger spread orientated parallel to strike. The apparent resistivity curve for the spread orientated normal to the dipping interface would produce a steeper curve, and be discontinuous as the electrodes crossed the surface interaction, see Kunetz (1966).

If the interface were buried, the apparent resistivity curve might show a slight irregularity, which would be considered as a lateral effect.

The interpreted results for all the V.E.S. are shown in drawing No. 74-458, and the pseudo-geological vertical sections are shown in drawing No. 74-451. The significance of the layers in the pseudo-sections is unknown at present, as no drill hole data is available in this area for a direct comparison of these layers with geological units or resistivity logs. It may be possible after a number of drill holes have been drilled and geophysically logged, to revise these interpretations based on the principle of equivalence (Kunetz, (1966)), to obtain a better understanding of the subsurface geology.

The interpretation procedures adopted for this area are as follows:-

- 1. To estimate the depth to fresh (unweathered basement), which is represented by the high resistive basement.
- 2. To estimate the ease which surface currents can flow above the infinite basement. This is represented by the Total Longitudinal Conductance of the layers above this basement. A thin conductor restricts the current flow, whereas thick conductors spread the current flow.
- 3. To estimate the possible depths to the top of the weathered basement and the variation in thickness of this horizon, and, to indicate conductive layers within this composite layer.
- 4. To estimate the conductance of the near surface saline or clay layers.

#### 7.1. RESISTIVITY BASEMENT

In terms of resistivity, the infinite resistivity basement can be considered as a well comented formation with approximately zero porosity and permeability to aqueous solutions. In the present, area of investigation, this may be considered geologically as unweathered 'Cleve Metamorphics' or possibly overlying Upper Proterozoic Rocks, i.e. Shelf Adelaidean.

of the 105 V.E.S. interpreted from the Port Broughton and Tickera data, 74% were interpreted as being terminated a high resistivity layer. This layer can be assumed relative to the previous layer to have a minimum resistivity contrast of 40 times, and have a thickness at least three times the depth to its top surface i.e. a transmissivity of 120 came, based on the two layer curves of Orellina and Mooney (1966). For the case stated, if the current electrode half distance AB/2 exceeds the above thickness ratio, then the apparent resistivity curve departs from the 45° asymptote, and becomes asymptotic to the true resistivity of that layer. In this survey, it was considered that for practical purposes half the current electrode AB/2 distance of the V.E.S. should not exceed 500 metres, as depths greater than these would not be egnomic for investigations of basement mineralization.

The interpreted final layer resistivity for the remaining 26% (i.e. 27) of the V.E.S. showed resistivities ranging from 4 to 760 ohm metres

The overall mean and standard deviation of these values were 187 and 184.2 ohm metres respectively. It was considered that the distribution represents two populations, the lower one, assumed to be less than 100 ohm metres, correlating with weathered basement and the other fresh basement. The parameters of these two groups are in Table 2 below.

TABLE 2
RESISTIVITY (IN OHM METRES)

RESISTIVITY BASEMENT TYPE	NO.OF INTERPRETED VALUES	RANGE	MEAN	STANDARD DEVIATION	RANGE FOR 849 CONFIDENCE LIMITS (mean + S.D.)
Weathered Basement	12	4 to 98	52	32	20 to 84
Fresh Basement	15	102 to 760	296	184	112 to 480

The purpose of the discussion of Table 2 was to estimate a possible range of resistivities for weathered basement, as there is no drill hole information available in this area.

#### 7. 1.2. CONTOURS OF RESISTIVITY BASEMENT

The contours of the depth to the resistive basement are shown in drawing No. 74-268. This is a composite, complied from the depth to the infinite resistive layer and the depth to the final finite resistive layers. The latter was assumed to be representative of the basement. The results show that the top of this horizon is very irregular, and may reflect either old topographic features or different rates of erosion, which in turn reflect differences

in basement lithology.

#### 7.2. TOTAL LONGITUDINAL CONDUCTANCE

The longitudinal conductance (S) in metres/ohm metre of any layer having a thickness Ei (in metres) and a resistivity Ri (in ohm metres) is defined as the ratio of the thickness to the resistivity of the ith. layer is given by Si = Ei/Ri.

Then the total conductance of a series of layers is given by the sum of these ratios.

In the case of an infinite resistive final layer in a series of horizontal layers, the rise of the curve is asymptotic to a 45 degree line. An estimate of the total conductance of all the layers can be determined by the ratio of the abscissa reading to the ordinate reading at the point on the line corresponding to the apparent resistivity value of unity.

(4 (医障。

The total conductance was estimated by the above method for the infinite layer case, and the total conductance for a finite resistivity last layer was determined as the sum of conductances using the interpreted results.

The contours of the total longitudinal conductance is shown in drawing No. 74-248. The areas of low overall conductance, located south of Tickera township and south-west of Alford, indicates thin faline areas of very shallow

basement. The area in the north, however, having total conductance ranging from 40 to 320 metres per ohm metre indicates a region of highly conductive material, which probably correlates with an area of salt water incursion. The remaining area indicates irregularities in total conductance values, and regional increase towards the north-west.

The interpreted anomalous frequency, I.P. effects, shown in drawing No. 73-708 suggested for further investigations occur in the areas of low conductivity, except for those located to the north and northwest of Alford. The latter is in the area of an older copper prospect, see drawing No. 73-709.

#### 7-3- WEATHERED RESISTIVITY BASISHENT

The resistivity values estimated above for the 84% confidence limit, Table 2 range from 20 to 84 ohm metres for an assumed normal distribution. These values compare with those interpreted from known shallow basement areas, south of Tickera township. Therefore, if the layer above the resistivity basement falls within the range greater than 20 ohm metres, it is assumed that this layer is weathered basement. However Genco drilling AH6, near VES 26 by Sibenaler (1974) intersected quartz, mica and weathered felspar? at 1.9 m, which indicated that the weathered basement could be compared with resistivity values as low as 15.5 ohm metres.

It must be noted, that geologically the top of the weathered basement is very ill defined, as most of the area is covered by a red clay or red sandy clay. Firman (pers. comm.), considered that this clay horizon in the Bute area, is purely an in-situ weathering product of the underlying basement, as it contains considerable relic features of the basement. It is probable that the same situation occurs in this area. The clay horizon has been sampled geochemically to define the geochemical basement, by J. Lynch, Chief Geologist of North Broken Hill Pty. Ltd. at Moonta. This geochemical basement occurs within this oxidized and weathered clay layer. The depth of weathering recorded in the drill holes at Kadina, extends to a depth of a few hundred metres. Therefore, using the following criteria a hypothesis for the weathered basement is tentatively defined.

- 1. The resistive weathered basement, has value greater than 20 ohm metres and generally occurs above the infinite basement.
- 2. In conductive environments, the final upturn or rising branch of the apparent resistivity curve generally with a slope of less than 45 degrees, was considered as a significant resistivity change, to represent the possible top of weathered basement. The shape of this curve may have been interpreted by either an 'A' or 'K' type curves.

The contours of the depth to the top of this interpreted weathered unit is shown in drawing No. 74-452. This in turn corresponds to the depth to the base of the low resistivity zone.

The area confined by the 30 metres contour can be considered as relatively shallow. The isopachs contours of

the interpreted thickness of this weathered zone are shown in drawing No. 74-457. The shaded areas in this plan are areas, in which the thickness of weathered zone is unknown, as it was not possible to resolve the weathered material in the V.E.S. The contours of the Traverse Resistance of the interpreted weathered layer shown in drawing No. 74-456, indicate areas of high resistance, i.e. grater than 1 000 ohm metre<sup>2</sup>, and the areas of low resistances, i.e. less than 250 ohm metre<sup>2</sup>. Some of these areas correlate with weathered basement clay, in the neighbourhood south of Alford.

#### 7.4.1. LOW RESISTIVITY LAYERS

The low resistivity layers were considered to be saline if the resistivities were less than 5 ohm metres. However the Gemco drilling, Sibenaler (1974) (DM 195/74) near V.E.S. Nos. 31 and in the neighbourhood of Alford, A.H. 1 near V.E.S. 66; AH 2 at VES 31; and AH 3, near VES 64 showed a thick sequence of red clays and sandy clays, which correspond to resistivities given below.

GEMCO DRILL	DEPTH OF DRILL HOLE	RESISTIVITY VALUES OF
HOLE NO.		CLAY (IN OHM METRES)
AH. 1	6.8 m	3.6 to 6.2
WH. 5	9.3 m	1.8 to 9.1

Results of the overburden rocks from an auger hole near AH1, drilled by North Broken Hill, Lynch (pers. comm.), showed that the clay extended down to approximately 30 metres, and by comparing this with the results of the first

30 m in VES 66, indicates that the clay can have resistivity of 3.66 ohm. Therefore as the clay has resistivities between 1.8 to 9.1 ohm metres, it is relatively impossible to distinguish between saline layers and clay in this resistivity range.

# 7.4.2. TOTAL CONDUCTANCE OF THE SALINE LAYER ABOVE THE INTERPRETED WEATHERED BASEMENT

The sum of the individual conductances of the layers above the interpreted weathered basement are shown in drawing No. 74-453. These contours indicate conductive rocks in the areas, east of Alford near Peela Weela Bore; in the vicinity of VES BP 13; and in the large region in the north of the survey area. The contours of the depth to the top of interpreted weathered basement drawing No. 74-452 shows that these areas already mentioned above, have thick saline deposits.

8. THE PEASIBILITY OF FUTURE INDUCED POLARIZATION AND ELECTROMAGNETIC METHODS BUING USED IN THE TICKERA AREA

The major problem with this area is the relatively thick cover of saline layers, highly conductive clays and the relatively deep high resistivity basement as can be seen in drawing Nos. 74-248 and 74-268. For the investigation of sulphides within the high resistivity basement, equipment will be required which can be used to separate the source signal from the induced coupling. This can be obtained by using the new time domain I.P. equipment, equivalent to the HcPhar Multi-Mode system, which uses a mean stacking process to obtain a representative phase values in the areas of high noise level. A number of phase-shifts can be recorded.

In the V.E.S. the AB/2 distance where the inductive coupling became significant was seen by the separation of the apparent resistivity curves for frequencies of 3 and 0.3 hz for each probe, and the half current electrode distance AB/2 (in metres) were recorded and compared with the interpreted depths results. The areas shown in drawing No. 74-455 indicate areas of strong electromagnetic coupling before the basement is reached. Therefore I.P. is not recommended to be undertaken in these areas.

#### 8.1. EXPECTED I.P. RESPONSE IN THE TICKERA AREA

The main type of mineralized zone in this region is a vien. Geophysically this can be represented by a thin vertical sheet of infinite vertical and lateral extent, i.e. dyke.

# EXPECTED SURFACE FREQUENCY EFFECT RESPONSE OF A MINERALISED DYKE.

Approximate frequency effect responses to outcropping and buried dykes can be predicted using data produced by Geoscience Inc. theoretical curves. The measured frequency effect depends on the following ratios:

- (1) Dipole spacing to dyke width.
- (2) Overburden resistivity to basement resistivity.
- (3) Overburden resistivity to dyke resistivity.
- (4) Dipole spacing to overburden thickness.
- and (5) the true frequency effect of the dyke.

It is thus necessary to specify these parameters in any representative model. Using the parameters applicable to the Tickera area; Table 3 has been constructed illustrating the resulting surface frequency effects with varying overburden, thickness and differing overburden resistivities.

The results are represented in terms of the dipole spacing (a) and at n = 4.

In the model, it is assumed that the resistivity ratio of the overburden to basement is 1:10, and that the dyke has a true frequency effect of 25%. The model and its parameters are shown in drawing No. S 10933.

The independent parameters (i.e. a, and n) are assumed to represent the range of valves, which would give optimum frequency effect valves.

The frequency effect is defined as the difference of resistivity measured at two frequencies, i.e. 3 and 0.3 hz. In this report the change in frequency response is given as the difference in frequency effect recorded for two different overburden resistivities with other factors remaining constant.

Table 3, shows that the change in frequency response decreases by 25% as the resistivity of the overburden decreases from 10 ohm metre to 5 ohm metre. Further extrapolation to say, 1 ohm metre overburden material is not considered to be justified as the frequency effect would be very small.

The former implies, that by using an I.P. dipole spacing approximately equal to the interpreted depth to the resistive basement based on the VES data; it is possible to resolve a thin vein of width a/3 at a depth equal to the dipole spacing at n = 4. Thinner bodies will not be resolved. For wider models, the I.P. response is larger and more persistent.

In the case of finite bodies having a square cross section 'a', where 'a' is the width and depth of body, the source is resolvable, if the overburden thickness is less than, or equal to, 'a'. If the overburden thickness is greater than the width of the body, then source will not be clearly resolved.

#### TABLE 3

Expected frequency effect for different resistivity surface layers.

#### EXPECTED FREQUENCY EFFECT

Overburden Thickness	Overburden Resistivity	10 ohm metre	5 ohm 1
8/3		8%	6%
æ		3 <b>.5</b> %	2.6%
2a		2%	1.5%

# 8.2. Electromagnetic Response in the Tickera Area

In considering, the electromagnetic response in the Tickera area, we can assume that the effective depth of penetration is twice the skin depth for a particular frequency.

The skin depth (s) equals 
$$\left\{\frac{2}{\text{Wu}}\right\}$$
 where R = resistivity (in ohm metres)

 $W = 2\pi f$ , where f is the frequency

$$u = 4\pi \times 10^{-7}$$

i.e. the skin depth S = 503.8 
$$\left(\frac{R}{f}\right)$$
 %

The presence of both saline and clay layers, representing conductive overburden have resistivities ranging between 1.8 to 9.1 ohm metres. The skin depth for the V.L.F. - IM frequency of 22.3 KHz ranges between 4.5 to 10.2 metres. In areas of lower resistivities, i.e. due to salt, having resistivities of less than 1 ohm metre the effective depth of penetration is 3.27 m. These depths of basement for VLF - FM frequencies is unsatisfactory, except in areas of resistive overburden.

The two other electromagnetic systems held by the section, are the McPhar Dual frequency R.E.M. system and the Turam E.M. system which operate at the following frequencies 1000 and 5000 hz, and 660 and 220 hz respectively. The skin depths for these frequencies for particular resistivities can be seen on (drawing No. 74-328) the monograph of skin depth versa frequency. It can be seen that the skin depth for the above measured frequencies for 1 ohm metre material are small compared with the skin depth estimate for the low frequencies of 3, 6.3. and 0.1 hz used for the V.E.S., which are 293 m, 928 m and 1607 m respectively.

If we consider using the Turam system, using both frequencies, the skin depth for 2 ohm metre material is 29 m and 50 m respectively i.e. the original signal has been reduced to 1/e of the amplitude i.e. 30% of the original amplitude. The depth of penetration for this low resistivity materials is at the most 60 to 100 m for the two specified frequencies. In this case the amplitude is only 10% of the criginal signal.

It is considered that the Turam system shouldbbe used

in areas, where the total conductance is less than 25 metres per chm metre for any success of this EM system using 220 hz. This area is shown in drawing No. 74-248. If 440 hz is all that can be obtained using 400 cycle I.P. supply generator. The areas to be investigated should have a total conductance of less than 18 metres per ohm metre.

# 9. <u>CONCLUSIONS</u>

The qualitative interpretation of the vertical magnetic intensity profiles has outlined possible areas of granite or acidic gneiss, other than the so-called Tickera Granite. It is interesting to note that an old copper prospect is reputed to be situated near Tickera township.

The re-appraisal of the original induced polarization data has revealed a number of additional frequency effects, which have been classified into a number of types, based on the correlation with apparent resistivity and magnetic data. A number of possibly significant I.P. anomalies for further verification with the new time domain equipment have been outlined, as shown in drawing No. 73-708.

The results of the feasibility study of future

Induced Polarization and electromagnetic surveys being undertaken in this area, show that both methods are appropriate,
but will have to be undertaken with considerable caution, and
in the case of electromagnetic surveys, using a low frequency
of the order of 440 hz or less. The areas for these future
surveys and areas of relatively shallow depth i.e. 60 metres

are outlined in drawing No. 74-455 together with the priority areas.

#### 10. RECOMMENDATIONS

Most of the future geophysical and geochemical surveys will have to be undertaken between the time, when the cereal crops are harvested and sown i.e. from January to May as most of the area is under cultivation. It is recommended that a number of auger holes using the Mayhew drill, be undertaken to obtain geochemical picture of the in-situ clays and samples of bedrock. These drill holes should be geophysically logged to obtain a better understanding of the interpreted resistivity layers. The geochemical data should provide additional target areas for geophysical investigation.

A number of areas where both Turam, and Input E.M. System, and Induced Folarization methods can be used, have been recommended for the future geophsical investigations and are outlined in Drawing No. 74-455.

- PRIORITY AREA 1 This area is relatively shallow and the
  a number of probable I.P. anomalies are
  present. In addition some unrecorded copper
  and gold prospects are said by the locals
  to exist in the Tickera township area.
- PRIORITY AREA 2 The interpreted area of possible granitic gneiss, based on the magnetic results and the location of some I.P. effects near the location of an old Copper prospect, make this area worthy of further investigation.

FRIGRITY ARDA 3 - This area contains a magnetic some type

1. which is comparable with the some
type 1 located south and southeast of
Alford, where promising secondary mineralization has been found by a company.

Those three areas recommended based on areas for best grouphysical results and need not represent the best areas for sineralization. All some type 1 and 2 areas, should be investaged geochemically especially the areas east of Alford, in areas of conductive overburden.

RAGANIS

ACQUATAME SENIOR GROSSYSICIST

AND ORATION GROSSICS SECTION

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  Parameters.

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#### APPENDIX 1

# PLRSCHAL AND SURVEY ROUTPHANT

PERSONNEL			DATE
R.A. GERDES	ACSISTANT SENIOR GEOPHYSICIST	6.4.73	to 19.5.73
	and	30.4.73	to 23.5.73
V.E. VIGHTMAN	Geophysicist	6.4.73	to 19.4.73
	and	30.4.73	to 11.5.73
L. CEST	TICHNICAL OFFICER	6.4.73	to 19.4.73
G. GAILBRATH	PIELD ACCIGTANT	9.4.73	to 19.4.73
P. TSANAKAS	FIELD ASSISTANT	30.4.73	to 23.5.73
C. LOSKOS	PIELD ACSISTANT	30.4.73	to 23.5.73

#### SURVEY EQUIPMENT

# Geophysical Equipment

Geoscience Induced Polarization Low Power Transmitter
Two McPhar P600/Induced Polarization Receivers
Reels of Wire (i) 4 x 500 metres

(ii) 2 x 50 metres

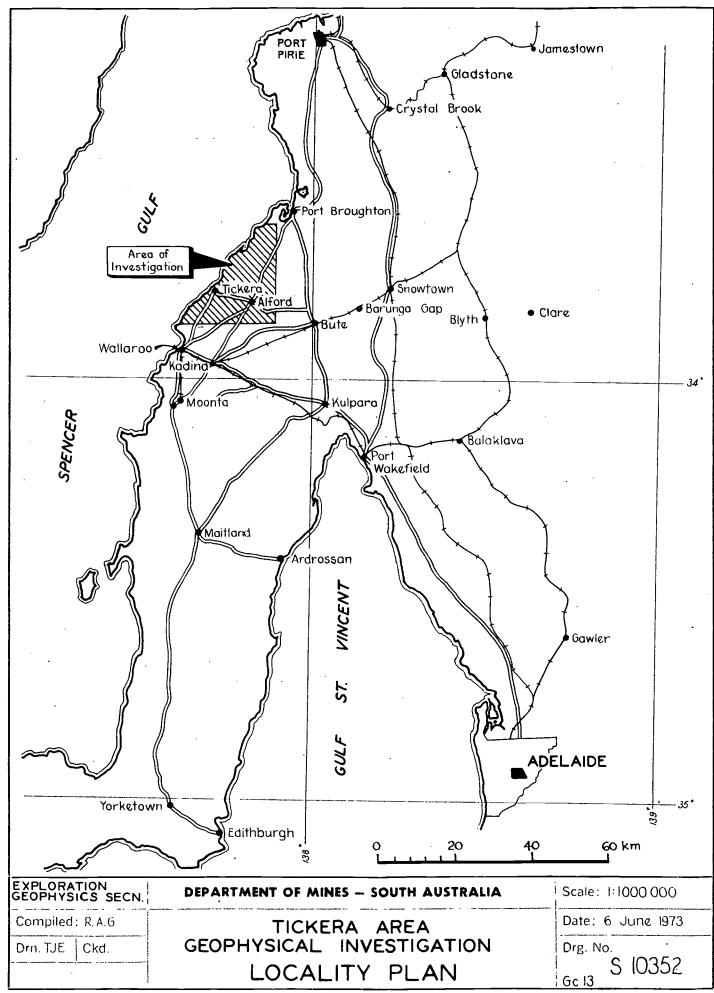
Stakes etc.

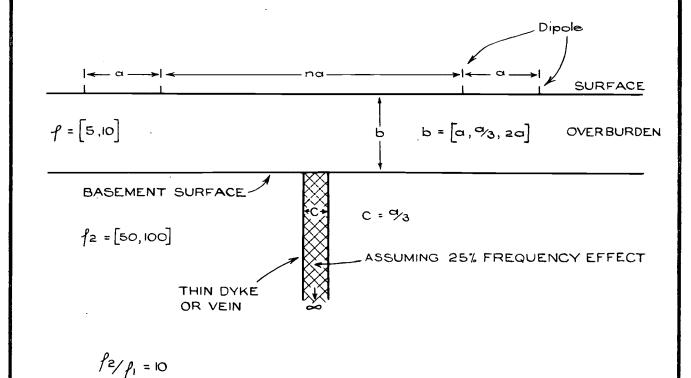
## Vehicles

Landrover DM. 168 with power take off (to power 115 volts 400 c/s 8000 generator)

DM. 25 Plat top .

Caravans DM. 39 and 73.





## LEGEND

a = Dipole spacing.b = Overbunden thickness taken on values a, 9/3 and 2a.

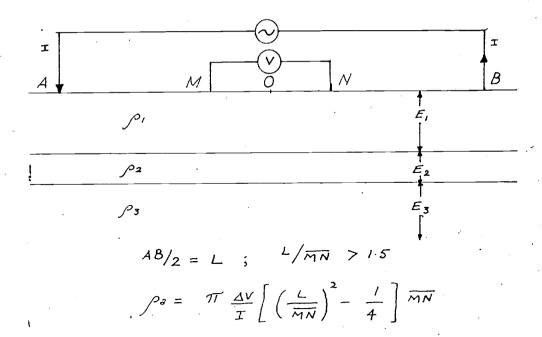
Width of dyke or vein. Has width %.

Resistivity of overburden taken on values

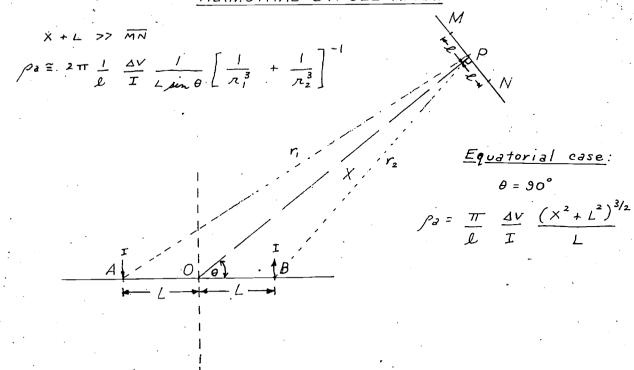
5 and 10 ohmmetres. Resistivity of basement taken on values 50 and 100 ohm metres.

DEPARTMENT OF MINES - SOUTH AUSTRALIA			
EXPLORATION	Drn.R.A.G.	TICKERA AREA GEOPHYSICAL SURVEY	SCALE:
GEOPHYSICS SECTION	Tcd. J.W.	DIAGRAM SHOWING THIN DYKE	S10933
R.A.GERDES	Ckd. AF	·	Gc 13
GEOLOGIST	Exd.	MODEL FOR INDUCED POLARIZATION	DATE: 31st July 1974

## SCHLUMBERGER ARRAY



## AZIMUTHAL DIPOLE ARRAY



DEPARTMENT OF MINES - SOUTH AUSTRALIA

Compiled: R, G, N.

Date: 26 - 7 - 7/.

Drn. Ckd.

Drg. No.

\$9564
Fd+1

