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TENEMENT:

S.M.L. 521

TENEMENT HOLDER: MINES EXPLORATION PTY. LTD

REPORTS:

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ASARCO (AUSTRALIA) PTY. LTD.

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ADELAIDE

SPECIAL MINING LEASE (PENNESHAW)

FINAL REPORT

bу

G.J. DREW Student Geologist

A.J. HOSKING Geologist

Adelaide, South Australia. 7 May, 1971. Technical report no. 33
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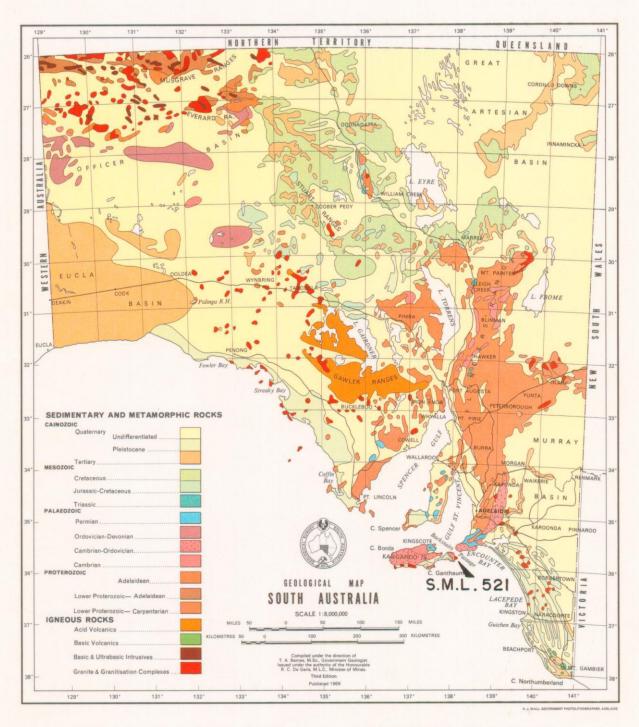
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SUMMARY

Geological investigations within S.M.L. 521 involved reconnaissance mapping together with stream sediment and rock chip sampling.

Most attention was directed to Lower Palaeozoic granitic rocks which display the effects of some hydrothermal alteration. Molybdenum and copper were the principal metals sought within these rocks.

Pyritic Cambrian metasediments were also prospected for base and precious metals.



LOCATION MAP, S.M.L. 521 PENNESHAW PROSPECT

Hydrothermal alteration zones in Lower Palaeozoic granitic rocks which occur on the eastern end of Kangaroo Island were considered to offer potential for base metal mineralization, in particular copper and molybdenum.

Pyritic horizons within Cambrian metasediments of the Kanmantoo Group were also thought prospective.

Special Mining Lease No. 521 covering the eastern end of Kangaroo Island was granted to ASARCO (Australia) Pty. Ltd. by the South Australian Department of Mines (see Fig. 1).

The results of the company's reconnaissance exploration program are contained in this report.

LOCATION AND ACCESS

Special Mining Lease 521 is situated approximately 70 miles from Adelaide (Penneshaw-Adelaide). The shortest direct route from the mainland to the island however lies between Penneshaw and Cape Jervis across the waters of Backstairs Passage, a distance of some 12 miles.

The lease is 90 square miles in area and covers approximately five percent of Kangaroo Island (see Fig. 2). The island itself is elongate in shape with maximum dimensions of 90 miles east-west and 35 miles north-south and an approximate area of 2,000 square miles.

The island is served by daily commercial air services from Adelaide to Kingscote and Adelaide to Penneshaw. A boat sails twice-weekly between Adelaide and Kingscote transporting vehicles, freight and passengers. This service is the principal means of supply for the island. A daily passenger service also exists between Cape Jervis and American River. In addition, vehicles and freight reach the island two or three times weekly by means of a boat service from Adelaide to American River.

The island is well covered by a network of public roads and access to most areas is relatively easy.

PHYSIOGRAPHY

The lease area is mainly a low, lateritic plateau with little overall variation in topography. The coastline however is rugged and in places precipitous cliff faces have formed, particularly in granite and aeolianite.

Most vegetation has been cleared from the flat-topped hills. Thick, low, mallee scrub remains in some areas and particularly in the south-east portion of the lease. Some stream courses are thickly covered.

The principal stream courses are the Willson River, Deep Creek and the drainage into Lashmar Lagoon-Chapman River. Numerous small creeks dissect the near-coastal areas. Stream courses for the most part are poorly defined.

Most of the approximate 23 inch average annual rainfall falls during the winter months.

Average seasonal temperatures tend to be appreciably lower than those experienced by Adelaide.

LEASE TENURE

ASARCO (Australia) Pty. Ltd. was granted S.M.L. 521 for a period of six (6) months commencing 10 December 1970. Two small areas were excluded from the lease, these being National Park Reserves. Also, all land extending one half mile inland from high water mark was excluded.

PREVIOUS EXPLORATION AND MINING ACTIVITY

Elcor Australia Pty. Ltd. held Special Mining

Leases 252 and 253 encompassing all of Kangaroo Island, with
the exception of the Flinders Chase Fauna & Flora Reserve.

Both leases expired on 31 October 1970. This company's work

(Ref. 2) was apparently confined to a near-coastal area west
of Kingscote where numerous, small, lode-type Pb-Zn deposits
occur (Ref. 1).

A narrow quartz-pyrite-arsenopyrite vein has been worked on a very small scale for Au in the past (Ref. 1).

Gem quality chalcedony (var. carnelian), beryl (var. aquamarine), tourmaline (pink, green and blue varieties) and garnet have been obtained from a large pegmatite. Clay

and feldspar have also been worked at this particular locality (Ref. 1).

A Mn occurrence is also known (Ref. 1).

Reports of Sn (Ref. 1) and native Cu (personal communication from a local farmer) occurrences ten miles southeast of Hog Bay and at Cape Hart respectively have not been confirmed.

METHODS OF INVESTIGATION

Reconnaissance Geological Mapping

The lease area was mapped at a scale of 1:50,000 using vertical aerial photographs and airphoto mosaics.

Aeromagnetic map coverage (Fig. 3) of the area was also utilized. Data was transferred to a base map compiled from published 1:50,000 topographic sheets. The results of this work are presented in Figure 4.

Lack of outcrop and dense scrub hindered the mapping. Most attention therefore, was directed to the continuous exposure along the coastline. It was hoped initially, to locate areas of interest on the coast and then extend investigations inland from these areas, beyond the half mile limit reserved from the lease.

Stream Sediment Sampling

Stream sediment sampling was carried out in conjunction with the reconnaissance mapping. The poor bedrock

outcrop, which is a result of the widespread laterite, calcrete and sand/soil covers in various portions of the lease, together with the poorly incised stream channels, meant that stream loads were rarely representative of the bedrock but rather of the surficial cover.

Most of the stream sampling program was concentrated upon the small creeks which drain directly into the sea and generally dissect bedrock, the Willson River and its tributaries and the drainages into Lashmar Lagoon.

A total of 146 stream samples were collected. The locations of these samples and their Cu, Pb, Zn and Mo values are shown in Figures 5 - 9. Table 2 in Appendix 1 summarizes the Cu, Pb, Zn and Mo contents of four size fractions of a number of samples. Table 3 in Appendix 2 summarizes Cu, Pb, Zn and Mo values for all samples collected.

Rock Chip Sampling

A number of rock chip samples were collected and analysed for various metals. The locations of these samples and their Cu, Pb, Zn and Mo contents are shown on Figures 5 - 9. Tables 4 - 6 in Appendix 2 summarize Cu, Pb, Zn and Mo values, additional metal values and a comparison of analytical techniques respectively.

Petrography

Petrographic descriptions of 15 rock samples were carried out by consultant petrologists. This data is contained in Appendix 3. The sample locations are indicated on Figure 5.

GEOLOGY

General

The regional geologic setting of S.M.L. 521 is shown in Figure 2 (Ref. 4).

The oldest non-igneous rocks exposed in the area are those of the Kanmantoo Group, these being metamorphosed Cambrian sediments. The metasediments have been intruded by granitic rocks in the vicinity of Cape Willoughby, by small basic dykes at False Cape and by pegmatites in several localities.

Permian sands, gravels and clays are reported from Penneshaw and leaf impressions of this age are recorded from a locality three miles ESE of Penneshaw (Ref. 4).

Early Tertiary bryozoal limestone is reported from Cape Willoughby. A small outcrop of late Cainozoic basalt occurs approximately three miles ESE of Fenneshaw. A petrographic description of this rock type is contained in Appendix 3, (ASARCO sample number A9923).

Pleistocene calcareous aeolianite occurs extensively in the southern portion of the lease. Sand dunes have been derived from this tock type and calcrete caps much of its exposure.

A thin veneer of lateritic ironstone gravel and podsolized soils covers much of the lease area.

Major Rock Units

Granite and associated rocks

Coarsely crystalline granite, adamellite and granodiorite occur in the vicinity of Cape Willoughby (henceforth in this report referred to collectively as granite). The granite contains distinctive "eyes" of bluegrey, opalescent quartz, biotite and both potash feldspar and plagioclase. Where exposed, the granite is usually fresh. Xenoliths of Kanmantoo Group metasediments are a conspicuous feature. Finer crystalline varieties of granite, aplitic dykes, pegmatites and occasional thin quartz veins exhibit cross-cutting relationships with the coarse granite. Tourmaline is a common constituent of many of the pegmatites.

Several types of hydrothermally altered rocks have been observed and rock sampling within the granite was concentrated upon alteration zones. In order of significance, muscovite, albite and epidote are the principal alteration In places, too, the introduction of potash during products. alteration has led to the development of distinctive, brickred Kspar. Greisenisation has resulted in the formation of quartz-muscovite rocks where alteration has been complete. Partially altered granite however, contains varying amounts of biotite and feldspar. The prominent "eyes" of quartz invariably remain in the muscovitized rocks. In extremely albitized rocks, the quartz "eyes" are no longer present and both the biotite and feldspar have undergone complete conversion. Epidote rims muscovite and albite alteration zones in some instances.

The larger exposures of altered granitic rocks lie to the south of Cape Willoughby. In particular, the headlands of Windmill Bay contain good exposures of altered rocks. On the northern headland, an east-west alteration zone approximately 150-200 feet wide is found. The actual volume of altered rock, both greisenised and albitised, within this interval however, is only approximately ten percent. Thin zones of altered rock, usually less than one foot wide, are separated by much larger zones of fresh, unaltered granite. A relatively structureless "pod" of mainly albitized rock, approximately 100 feet in diameter, occurs on the southern headland.

The distribution of the altered zones is largely controlled by a near-vertical joint pattern in the granite.

Other joint sets also have associated alteration effects.

Several very thin (less than one foot) alteration zones occur between Cape Willoughby and Pink Bay, but are lacking between Pink Bay and the granite margin, which is exposed further west on the north coast.

The contacts of the granite and the Kanmantoo Group metasediments are sharply defined. The granite near the contacts is strongly fractured and contact metamorphic features are not pronounced. Numerous and frequently large xenoliths of country rock included in the granite near the contacts, are prominent. In many such xenoliths, blue opalescent quartz and feldspar crystals are seen. Indications are that the granite was intruded as a relatively cool and substantially solid mass. The distinctive aeromagnetic anomaly

(Fig. 3) which trends southwest from Cape St. Albans may indicate a partially faulted contact between the granite and the Kanmantoo Group.

Pegmatites, which dissect both the granite and the Kanmantoo Group, contain coarsely crystalline quartz, muscovite and feldspar. Tourmaline is also an extremely common constituent of these bodies.

Kanmantoo Group metasediments

The Kanmantoo Group within S.M.L. 521 consists largely of quartzites, hornfels, phyllites and schists. Throughout much of the lease area the rocks strike approximately northeast-southwest and dip to the southeast — at 50°-80°. Sedimentary facings and cleavage-bedding relations indicate that the gross structure of the metasedimentary sequence is an overturned anticline. Individual units and detailed structure within the Kanmantoo Group were not mapped, however.

The rocks generally belong to the greenschist facies of regional metamorphism. In addition, metasomatic effects are evident, particularly near Cuttlefish Bay, where most rock sampling of the Kanmantoo Group was concentrated. Here, biotite schists and hornfelses containing actinolite and scapolite occur, together with minor tremolitic marble bands. Calcite-actinolite vein rocks are also found, often highly contorted.

Epidote, zoisite, chlorite and sphene are common minerals in the metasomatized schists and garnet has also been observed. Cordierite and andalusite have been described

from one specimen.

Basic dyke rocks

A number of finely-medium crystalline dykes of gabbroic composition occur at False Cape where they intrude Kanmantoo Group quartzites. The dykes, which vary in thickness from four to 25 feet, are perpendicular to the stratification of the quartzites and have sharp contacts with them.

MINERALIZATION

Stream Sediment Sampling

A summary of analyses of the 146 stream sediment samples collected is presented in Table 1. Full details are supplied in Appendix 1.

Variation of Mo content with particle size was observed (Table 2, Appendix 1). Consequently, the -20+40 fraction was utilized for analysis of all samples collected. Molybdenum was not detected in most of the -80 mesh size fractions analysed.

Table 1

<u>Distribution of metal values</u> <u>in stream sediment samples</u>

Cu	Value p.p.m.	- 5	5	10	15	20	25	40		
	No. of samples	52	51	19	15	4	4	1		
Pb	Value p.p.m.	- 5	5	10	15	20	25	30	35	40
	No. of samples	4	25	29	27	14	14	8	7	5

Pb	Value p.p.m. No. of samples				, .	018
Zn	Value p.p.m. No. of samples					
	Value p.p.m. No. of samples			70 1		
Мо	Value p.p.m. No. of samples					12 1

The results are characterised by low background values which are due to dilution of the drainages by surficial cover. Previous stream sampling by Elcor Australia Pty. Ltd. on the northern side of the island also indicated a similar situation in areas of cover in that region (Ref. 2).

Threshold values of 25, 40, 50 and 6 p.p.m. can be assigned for Cu, Pb, Zn and Mo respectively. However, samples with metal values exceeding these figures and which might be considered anomalous, do not indicate the presence of mineralisation. Rather, they appear to be random higher figures, and not significant departures from the distribution patterns.

Rock Chip Sampling

The results of rock chip sampling are summarised in Tables 4 - 6, Appendix 2.

Altered rocks within the granite have very low metal contents. In particular, analyses for Cu and Mo were disappointing and no evidence of leaching of appreciable quantities of sulphides was observed. Minor disseminated

limonite was found in places but is due to the breakdown of pyrite. Fresh pyrite also occurs, sparsely distributed in the granite and the altered rocks. In one sample only, traces of chalcopyrite were seen and this particular sample returned a Cu value of 560 p.p.m. No molybdenite was found in any of the thin quartz veinlets. Values of 18 and 8 p.p.m. Mo were obtained from two samples of one such veinlet in which pyrite was present.

Several higher Pb values, and in particular one of 2400 p.p.m., could not be accounted for. This particular sample was collected across an alteration zone six inches wide. The presence of Pb in greisenised and albitised rocks is unusual and the values may be the result of analytical error.

Tungsten and Sn, which are often associated with hydrothermally altered granitic rocks, were not determined in significant amounts while Ag and Au were not detected.

The base metal contents of the Kanmantoo Group rock types sampled in the lease area are low. Pyritic phyllites, hornfelses and schists were of principal interest.

Pyrite generally is associated with rocks which show the effects of some metasomatism. However, sampling, particularly near Cuttlefish Bay, did not indicate significant base metal concentrations in these rocks. In addition, Ag and Au were not detected. A thin quartz-pyrite-arsenopyrite vein which averages several inches in thickness and occurs at the margin of a pyritic band, gave values of 0.04ozs. per long ton Au, 0.25% Cu and 27.3% As. Arsenic values to 0.4% were found in

other samples of the pyritic rocks. One sample containing 0.15% Cu was collected from a thin calcite-actinolite vein. Molybdenum was not detected in the majority of samples, the highest value obtained being 22 p.p.m.

The thin basic dykes at False Cape have no associated base metal mineralisation, although minor pyrite does occur.

A narrow quartz-pyrite-arsenopyrite vein which was worked in a minor fashion for Au prior to 1933 (Ref. 3) contains no metal values of note.

CONCLUSIONS

Altered zones in the Lower Palaeozoic granitic rocks of S.M.L. 521 are mostly extremely small and unmineralised.

Pyritic horizons in the Kanmantoo Group within the lease also lack significant metallic mineralisation.

Narrow quartz-pyrite veins offer no potential for further prospecting, for the above reasons.

RECOMMENDATION

It is recommended that S.M.L. 521 be relinquished.

A.J. HOSKING Geologist.

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APPENDIX 1

STREAM SEDIMENT SAMPLING RESULTS

STREAM SEDIMENT SAMPLING RESULTS

Introduction

The results of analyses of stream sediment samples collected within S.M.L. 521 are presented in Tables 2 and 3.

All analyses were carried out by atomic absorption spectroscopy at the Australian Mineral Development Laboratories, Flemington Street, Frewville, South Australia.

All samples were analysed for Cu, Pb, Zn and Mo, the lower limits of detection being 5, 5, 5 and 3 p.p.m. respectively.

Most samples required drying prior to sieving. The -80 mesh size fraction was utilised in the case of Cu, Pb and 2n and the -20+40 fraction for Mo.

All results are in p.p.m. unless otherwise indicated.

Results

Table 2
Comparison of metal contents of four size fractions

	-20+40	-40+60	-60+80	-80
Sample No.	Cu Pb Zn Mo	Cu Pb Zn Mo	Cu Pb Zn Mo	Cu Pb Zn Mo
Penneshaw			Sx.	
A 8802	5 10 20 3	- 5 10 15 3	- 5 10 20 3	- 5 15 40 3
13	-5 10 5 3	" 10 10 -3	" 5 10 3	" 15 10 - 3
34	5 35 10 3	5 20 10 3	10 15 15 5	15 20 35 "
45	5 50 15 3	-5 25 15 11	10 25 15 6	10 35 35 "
50	15 20 20 7	5 -5 10 -3	5 - 5 5 4	-5 5 5 "
53	5 20 15 5	5 5 10 "	5 " 5 -3	5 5 10 4
54	25 10 70 4	15 5 55 "	10 5 40 "	15 5 50 "

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Sample No.	-20+40	-40+60	-60+80	-80
	Cu Pb Zn Mo	Cu Pb Zn Mo	Cu Pb Zn Mo	Cu Pb Zn Mo
Penneshaw 8858 59 62 70 75 87 8895 10803 7 10824 10906 16 36 46 55 67 73 93 A10997	5 20 20 4 5 20 20 -3 15 10 25 -3 15 10 15 15 5 40 15 -3 5 40 15 20 15 15 15 30 5 15 15 30 5 15 15 30 5 15 15 30 5 10 40 30 10 5 5 20 10 5 5 20 10 5 5 20 10 25 -3 20 30 10 30 10 5 5 20 10 5 5 30 10 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	5 5 15 15 15 15 15 15 15 15 15 15 15 15	5 10 -3 " 4 3 10 5 10 10 10 10 10 10 10 10 10 10 10 10 10	5 5 20 -3 5 10 15 " 10 10 25 " 10 10 25 " 10 15 25 15 " 20 25 70 " 20 25 70 " 15 10 10 3 10 20 35 " 10 10 25 " 10 10 45 "

Table 3

Minus 80 mesh Cu,	Pb, Zn	and -20+40	mesh Mo	values
Sample No.	Cu	Pb	Zn	Mo
A8801	- 5	10	45	- 3
2	11	20	30 20	3
) 1	11	15	20	5
1 5	11	40 15	50 🗸 15	 3
2 3 4 5 13 15	ff:	10	 5	フ 3
	11	10	ĩi	3
21 30	11	10	11	3
30 31	11: 11	40 45	20	8
34	5	45 35	1 5 15	8
35	5	40	15	ラ 3
36	5	40 70	15	5
37	5	30	15	. 8
38 39	5	50 45	15	3
40	う 5	42 50	20° 10	6
41	15	20		6
44	5	65	25 20	5
45	5	60	20	3
46 47	<u>ხ</u>	50 20 65/ 60/ 65/	15	3
41 48	う 5	65° 40	20 15	3
48 50	55555555555555555555555555555555555555	5	5	35353338833583666533337
	/	- 2	-	

	10	7 &		
			م	052
Sample No.	Cu	Pb	Zn	Mo
8853 54	5′ 15′	5 5	10 50	5 4
58 50	5	5 5 10 10	20	4
62	10	10	25	5 4 -3 -4 -3
67 68	5 5	10 5	15 15	-3
54 58 59 62 67 68 69 70	10 5	10 5 15 25 10	25 20	13°
71	15	10 15	35	11
72 73 74 75	5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	10	50 15 15 15 15 20 10 10 20 42 25 50 10 20 20 20 20 20 20 20 20 20 20 20 20 20	
74 75	5	10 10	20	- 5
79 81 . 82 83	1 <u>5</u> 10	15 20	45 ~ 20	3 8
. 82 83	10	15 15	25 25	3
84 85	5	15	25	- 3
86	10 5 5 5 10 10	15 15 15 15 30 25 30	20	6
87 88	10 5	25 30	30 20	<i>3</i>
84 85 86 87 88 89 90	10 5	10 20	10	- 3 5
91 92 93 95 96 97 98 8899	5 10 5 10 5 5 20 10	45 10	20 20 15 25 70 20 25 25 15	87638353563335553333333
93 95	5	20	25 70 er	3 -3
96 96	10	20 20 10 30 65 15	20	3
98	- 5	65 /	25 25	- 2
8899 108 01	11 15	15 10	15 45	11
2	5 5	5 5	15 15	#: 5
4	ii 5	5	25 45	5 3 11
6	15	5	35	11: °
8	10	5	25	11
9 10	5 15	10 5	25 35 /	11) 13
11 12	5) 25	5 10	35 / 45 /	5 -3
20 21	10	5	35 -))
22	- 5	25	15	tti M
24 27	5	10	5	3
10801 2 3 4 5 6 7 8 9 10 11 12 20 21 22 24 27 31 32 10833	1555 550055 555555555555555555555555555	10 55 55 55 10 50 55 55 50 55 50 55 50 55 50 50 50 50	45555555555555555555555555555555555555	3 -3
10833	15	20	35 🖑	18

Sample No.	Cu	Pb	Zn	Мо
109 0234568901234567892333456789012564789012345666666667890 1111111111223333333333444444455555555666666667890	555555550000550005" 55" "" 14-2-" 155" 15505" "" 17-1-" 17" 17-" 17	1221 - 131211115110055050505055555555000555505050505050505	20055005555050505555555555555555555555	3"373 3343333456563643663665552643"55555353"353""353""33

Sample No.	Cu	Pb	Zn	Мо	027
Sample No. 10971 72 73 75 79 80 81 87 90 91 92 93 94 95 97	0 - " 5555555555555555555555555555555555	15055555055055505	Zn 2555555555555555555555555555555555555	Mo 353633331 1 3333336	
A10999	'n	5	15	<u> </u>	

APPENDIX 2

ROCK CHIP SAMPLING RESULTS

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ROCK CHIP SAMPLING RESULTS

Introduction

The results of analyses of rock chip samples collected within S.M.L. 521 are presented in Tables 4, 5 and 6.

Analyses, with the exception of five (5) Au determinations, were carried out at the Australian Mineral Development Laboratories, Flemington Street, Frewville, South Australia. A number of analytical techniques were utilised. The five (5) Au analyses were carried out using atomic absorption spectroscopy by Mc.Phar Geophysics Pty. Ltd., 50 Mary Street, Unley, South Australia. The particular method used had a lower detection limit of 0.5 p.p.m. (value suffixed M).

All samples were analysed for Gu, Pb, Zn and Mo. Selected samples were analysed for Ag, Au, W, Sn, Co, As. The lower limits of detection for the particular methods used were (in p.p.m.) Cu (5), Pb (5), Zn (5), Mo (3), Ag (1), Au (3 - also 0.01 ozs./long ton), W (50), Sn (1), Co (5), As (10).

Gold analyses additional to the above were carried out by semi-quantitative emission spectroscopy (samples in ranges A8806 - A8866, A10814 - A10845).and fire assay (samples in range A9901 - A9923, reported in ozs./long ton and indicated thus *).

Copper, Pb, Zn, Mo and Co were determined by atomic absorption spectroscopy, and W and Sn by semi-quantitative emission spectroscopy. Arsenic was determined by a modified

Gutzheit method.

All results are in p.p.m. unless otherwise indicated.

Results

Table 4
Cu, Pb, Zn. Mo values

Sample No. A 2298 99 2300 6229 30 31 32 33 34 35 36 37 38 39 40 41 42 43 44 57 Not 58 59 60 6261 8806 7 8 9 10 11 12 14 16 18 19	Cu 50055555555 11 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 15 16 15 17 15 18 15 19 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15 10 15	Pb 120 2400 75 1400 850 250 6 280 180 100 100 100 100 100 100 100 100 1	Zn 25505550505000500550055005555555555555	Mo - 3
14 16 18 19 20 22 23 24 25	¹ 5	20 10 5 -5 10 10 10	15 55 55 55 -5 -5 -5	11.

,			prop.	032
Sample No.	Cu	Pb	Zn	Mo
10837 38	40 70	15 35 30	70 35	-) 5
10837 38 39 40	30 100	30 10	70 35 100 30 55	3 3
41	50	15	55 10	3
41 42 43	30 100 50 -5 190	25 3 5	50	3
44 10845	15 20	15 15	10 30	3 3
10901	45 45	10 15 35 35 15 15 10	30 60 85	75333533333"333333"33333"
. 7 -20	41-1 42246 424 	15 20	-5	- 3
21 24 25 26 27 28	15 5	20 10	120 5	3
25	5	140 10	5	3
27 27	45	10	15	3
28 29	25 - 25	10 15	25 25	3
29 30 31 32	45 65	5 35	12 12 12 12 12 12 12 10 10 10 10 10 10 10 10 10 10 10 10 10	3 - 3
32	5	5	55 400	3
44 56 58 60	45 25	10	65	, <u>-</u> 3
58 60	45 5	20 10	55 5	.11
74	5	25	65	
74 76 77 78 82 83		5 5		" 11:
78 82	#. 5	5 10	5 15	. 4
83	- 5	5	10	- 3
84 85	5 -5 -5 -5 120 -5	5 5	- 5	, tt
85 86 88	120 - 5	5 5	15 50	-3 3
89	íi 5	10 10 10 10 10 10 10 10 10 10 10 10 10 1	5 10 5 5 5 15 15 65	3 -3 3 -3
A 11000	9	ý	U)	-)

Table 5

Other metal values

Sample No.	Ag	Au	\mathbf{W}	Sn	Co	As
A 2298	-1					
99	11					
2300	11					
6229	11				•	
30	11					
31	II	-0.5M				
30	11					

Sample No.	Ag	Λu	W	Sn	Co	As	033
62§3 34 35 36 37 38 39 40	-1 "" "	-0.5M					
38 39 40 41 42 43	11 11 11 11 11	, II					
44 57 58 59	†† †† †† ††	H ~					
6261 8806 7 8 9	11 11 11 11	-3 ""	-50	5 3 3 2 15 1			
11 12 14 16 -18	# # # # #	11 11 11 11 11	11 11 11 11 11 11	15 1 1 2		,	
19 20 22 23 24 25 26	11 11 11 11	11 11 11 11	11 11 11 11	1 2 1 2 5 5 10 1 1			
25 26 27 28 29	11 11 11 11	17 11 11 11	11 11 11 11	1 5 10 1			
27 28 29 23 32 43 45 55 55 56 66 66 66 88 88	11 11 11 11	11 11 11 11	11 11 11 11	1 1 5 1			
51 52 55 56 57	-1 !!	11 11 11 11	11 11 11 11	.13151351			
60 61 63 64 65	11 11 11	11 11 11 11	.11 11: 11 11	-3 5 1 -1			
8866	II:	11	11	1			

Sample No.	Ag	Au	M	Sn	Co	As
9901	1	0.04*			2500	27.3%
9901 2 3 4 5 6 7 8 9 9923	1	-0.01*			40 30	4000
) 1	11	tt			40	140 830
5	11	11			30	100
6	tt	11:			30	. 60
7	.11. 11	18 11			35 20	15 30
8 9	n	n			20 35	25
9923	11	11			35 35	30 25 90
10814	11	3	- 50	2	- 10	
15		11	19	ე ე		
15 16 17	11	11	11	2 5 2 5 1		
18	11	11	11	5		
19	11°.	t† 11	11 11			
23	11	¥1	11	1 10		
18 19 23 25 26 28	tt	1.1	ŧì	15 20		
28	11	ŧŧ	ii	20		
29	.11 11	11 11	11	2		
30 34 35 36 37	;; ;t	11	u	2 2 2 2 1		
35	**	.11	11	2		
36	tt 	tt	11			
37	†† ††	t i	31 11	1 _1		
20. 39	ti	11	11	- ₁		
38 39 40	tt	11	n	-1		
41	11	11 .tt	11	1		
42	# #	.11	11	-1		
41 42 43 44	11	11	11	5 1 3		
A 10845	11	11	19	3		

Note: M - Mc.Phar determination
* - ozs./long ton

Table 6
Comparison of analytical methods

Sample No.	Cu 1 2	Ръ 1 2	Z n 1 2	Mo 1 2	Ag 1 2
	(5)(0 _* 5)	(5) (1)	(5) (20)	(3)(3)	(1) (0.1)
A 10814 15 16 17 18 19 25 26 29 34 37 39 41 42 43 44 44	35 30 30 30 30 30 30 30 30 30 30	35 100 50 100 30 200 150 200 150 20 10 100 5 5 10 100 5 1	120 100 190 180 130 90 220 200 270 200 100 100 60 50 10 -20 5 " " 90 95 80 20 -20 70 20 55 -20 50 -20 50 -20 50 -20	10503033""""""""""""""""""""""""""""""""	-1 0.2
29 30 34 35 36 37 38 39 40 41 42 43	20 20 5 10 25 30 10 40 55 30 40 25 70 50 30 25 100 90 50 40 -5 20 190 120	5 10 30 50 5 5 15 20 15 10 35 20 30 50 10 8 35 30 35 40	95 90 90 80 20 -20 70 40 35 20 100 120 30 20 55 60 10 -20 50 60	11 11 11 11 11 11 11 11 11 11 11 11 11	# 0. # 0. # 0. # 0. # 0. # 0. # 0. # 0.

Note: 1 - AMDEL atomic absorption spectroscopy

^{2 -} AMDEL semi-quantitative emission spectroscopy

Detection limits in brackets

APPENDIX 3

PETROGRAPHY OF 15 ROCK SAMPLES

Introduction

Specimens were described by H.W. Fander (1-13 inclusive) and I.F. Scott (14-15 inclusive) of Central Mineralogical Services, 231 Magill Road, Maylands, South Australia.

Thin sections are stored at ASARCO (Australia) Pty. Ltd., 323 Wakefield Street, Adelaide, South Australia.

Petrography

1. ASARCO sample no. A9902: C.M.S. section no. 4540

Identification: Metasomatised quartz-mica schist

Hand specimen: Dark, fine-grained, schistose rock with fine,

disseminated sulphides.

Microscopic: This is an extensively scapolitised quartz-mica schist.

The original rock, still present as layers and lenses, consists of very fine-grained quartz and parallel flakes of biotite and minor muscovite. Throughout the rock, ovoid poikiloblastic patches of scapolite have formed as a result of metasomatism involving replacement of most of the quartz, with survival of the small biotite flakes. Poikiloblastic patches of actinolite are also common. Granular sphene is present.

The sulphides are almost certainly associated with the metasomatic minerals and with late-stage carbonate-chlorite veins.

2. ASARCO sample no. A9906: C.M.S. section no. 4541

Identification: Metasomatised quartz-biotite schist

Hand specimen: Dark, fine-grained schist with layers or parallel veins of sulphides.

Microscopic: Very similar to A9902 in origin and history and may be termed a metasomatised <u>quartz-biotite schist</u>. It is more distinctly layered than A9902 and with a more obvious association of metasomatic minerals and the sulphide.

The rock consists of contorted layers of very fine quartz biotite schist and poikiloblastic scapolite and layers of coarsely crystalline scapolite, quartz, actinolite, biotite, sphene and opaques (sulphide).

3. ASARCO sample no. A9907: C.M.S. section no. 4542

Identification: <u>Impure marble (tremolitic)</u>

Hand specimen: Pale crystalline rock with ?relicts of schist.

Microscopic: This consists mainly of coarsely crystalline carbonate (not calcite; probably dolomite) and may be termed an impure marble.

The interlocking crystals of carbonate contain sheaves of acicular tremolite, some actinolite flakes of pale biotite or phlogopite, patches of quartz and sulphides. The carbonate crystals show strain extinction due to stress.

Since tremolite is the most abundant silicate, this may be termed a tremolite marble.

4. ASARCO sample no. A9908: C.M.S. section no. 4543

Identification: Actinolite - scapolite rock (metasomatic)

Hand specimen: Medium-grained, green ?amphibolite.

Microscopic: An extensively metasomatized rock; although

it contains abundant amphibole, it is not an amphibolite in

the normal sense. Very little of the original rock remains; it appears to consist of quartz and feldspar (?albite).

Very abundant, prismatic, poikiloblastic actinolite has developed throughout, with random to sub-parallel orientation. Lesser amounts of scapolite, clinozoisite-epidote and sphene also occur. These are all metasomatic, i.e. replacive. They do not appear to be accompanied by sulphides in this rock.

Minor carbonate also occurs.

- 5. ASARCO sample no. A9910: C.M.S. section no. 4544 Identification: Quartz - biotite - garnet schist Hand specimen: Dark, fine-grained, folded biotite schist Microscopic: This is a folded quartz - biotite - garnet It consists of fine mosaic quartz and small parallel schist. flakes of brown, biotite. Occasional euhedral porphyroblastic garnet crystals occur sporadically, and there is a garnet-rich layer adjacent to a folded quartz layer. The folding is a post-metamorphic phenomenon, with introduction of quartz-biotite veins or layers cutting across the schistosity; these layers have selvages of recrystallised biotite and may have formed by "lateral secretion" in an immediately post-metamorphic phase. The rock is a metasediment belonging to the greenschist facies of regional metamorphism.
- 6. ASARCO sample no. A9912: C.M.S. section no. 4545

 Identification: Quartz biotite garnet schist

 Hand specimen: Dark, contorted, fine-grained biotite schist.

 Microscopic: A quartz biotite garnet schist, very

 similar to A9910, with only minor differences.

The original sediment must have contained sand-size quartz grains as well as fine quartz and clays. These coarser grains have survived recrystallisation and occur as relict grains embedded in the schistose biotite-quartz matrix. Euhedral, poikiloblastic garnet crystals are sporadically distributed through the rock.

7. ASARCO sample no. A9915: C.M.S. section no. A4546 Identification: Garnet - zoisite - biotite - quartz schist Hand specimen: Dark, quartz - biotite - garnet schist. This schist is generally similar to the Microscopic: previous ones but there are mineralogical differences. In addition to quartz, biotite and garnet, zoisite and chlorite are conspicuous. Also, there is more garnet. Areas of poikiloblastic actinolite occur and there is sphene and carbonate. Garnet porphyroblasts are set in a matrix of orientated biotite. mosaic quartz and granular zoisite. A few layers or elongate lenses contain metasomatic carbonate, actinolite and minor opaques (fine suphides). The rock belongs to the greenschist facies.

8. ASARCO sample no. A9916: C.M.S. section no. 4547

Identification: Scapolitised biotite schist.

Hand specimen: Dark, knotted schist.

Microscopic: This is a scapolitised <u>quartz - biotite - actinolite schist</u>.

The rock consists of porphyroblasts ("knots") of pale, actinolitic amphibole in a fine matrix of quartz and lineated biotite. The actinolite evidently formed contemporaneously with metamorphism. The scapolite however, which also occurs as

poikiloblastic patches, is a post metamorphic phase since it contains orientated inclusions of biotite parallel to the schistosity. Fine opaques are scattered through the rock.

9. ASARCO sample no. A9917: C.M.S. section no. 4548

Identification: Epidote amphibolite

Hand specimen: Coarsely - crystalline, dark-green, amphibolite

with epidote.

Microscopic: May be termed an <u>epidote amphibolite</u> and containing small patches of scapolite.

The epidote occurs as well-formed, small prismatic crystals showing pronounced parallel alignment. The pale hornblende is the major mineral as large, interlocking patches with embedded aligned epidote. Small flakes of green biotite are present, especially in epidote-rich streaks.

Textural relationships suggest that the rock was a scapolitised biotite-epidote schist which was subsequently metasomatised forming the hornblende. Hence, although hornblende is the main mineral, this is not an amphibolite in the normal sense.

10. ASARCO sample no. A9920: C.M.S. section no. 4549

Identification: Quartz - biotite schist with scapolite.

Hand specimen: Finely-laminated, spotted schist.

Microscopic: A fine-grained <u>quartz - biotite schist</u> with small, ovoid "spots" or poikiloblasts of scapolite.

The finely laminated fabric of the rock is due to alternating biotitic and more quartzose layers. The scapolite poikiloblasts are late or post-metamorphic and contain orientated biotite flakes; they are often surrounded by a halo of fine, granular sphene.

A transgressive, quartz-carbonate vein occurs, cutting across the schistosity at a low angle but lacking sharp borders. Detrital heavy minerals are seen e.g. green tourmaline and zircon.

11. ASARCO sample no. A9922: C.M.S. section no. 4550

Identification: Schistose, tremolite marble

Hand specimen: Well-crystallised, streaky, brown mica schist.

Microscopic: This is actually an impure schistose marble.

It consists of elongate, parallel grains of carbonate (probably dolomite) with sub-parallel, large, prismatic crystals of tremolite with carbonate inclusions. Wisps of phlogopite are very common throughout, as well as thin flakes of colourless chlorite. Small grains of sulphide (?pyrite) occur sporadically. The paler streaks in the rock are caused by thin layers containing little or no phlogopite.

12. ASARCO sample no. A9923: C.M.S. section no. 4551

Identification: Basalt

Hand specimen: Very fine-grained, dark-grey, igneous rock.

Microscopic: This is a slightly vesicular, porphyritic

basalt. It is composed of occasional phenocrysts of augite and microphenocrysts of labradorite set in a groundmass of small labradorite laths, granular pyroxene, opaques and brown glass. The small, spherical vesicles are filled with pale chlorite and yellow chlorite occurs in the groundmass, generally adjacent to the vesicles. However, on the whole the rock is quite fresh.

13. ASARCO sample no. A9924: C.M.S. section no. 4552

Identification: Scapolite - epidote - actinolite rock

Hand specimen: Pale, crystalline rock with amphibolitic

streaks.

Microscopic: A scapolite - epidote - actinolite rock.

The principal mineral is scapolite as large, interlocking anhedral plated containing sub-parallel lines of small, lineated epidote crystals and large poikiloblastic crystals of actinolite. This mineral also occurs as radiating fibrous groups. Granular sphene is conspicuous in some parts and pale chlorite flakes are seen, interstitial carbonate is rare.

The larger scapolite crystals contain small biotite flakes and the original host rock may have been a biotite schist, now almost completely metasomatised.

14. ASARCO sample no. A10924: C.M.S. section no. 4915

Identification: Pyroxene hornfels assemblage

Hand specimen: A pale-grey, fine-grained, metamorphic rock adjacent to a white buff siliceous portion.

Microscopic: The light-grey portion of the rock contains quartz, potash feldspar and diopside mainly. For the most part, the rock is medium-grained and average grain size is less than 0.5 mm. The texture is granular and rather horn-felsic. Poikiloblasts of diopside occasionally mach dimensions greater than 1 mm. Accessory sphene is present. The rock is a member of the <u>pyrexene hornfels</u> facies (orthoclase is present rather than microcline) and its composition suggests that the primary rock was a calcareous, feldspathic sandstone.

15. ASARCO sample no. A10944. C.M.S. section no. 4917

Identification: Biotite - quartz schist with

poikiloblastic cordierite and andalusite

Hand specimen: Dark-brown biotite - andalusite schist.

Microscopic: This rock is essentially a medium-grained biotite - quartz schist with numerous purphyroblasts developed within it. At least two varieties of porphyroblasts are present namely cordierite with typical quartz and also random mica (muscovite and biotite) inclusions and also andalusite which is also full of similar inclusions. Perhaps both minerals should be described as poikiloblastic.

It appears that the andalusite was first formed and has been partially deformed by the metamorphism. Cordierite has formed as rims on the andalusite or as separate porphyroblasts of a slightly younger age. There is a noticeable presence of minor muscovite replacing biotite adjacent to cordierite as well as inclusions within the cordierite. It appears that magnesium and potash metasomatism of a previously regionally metamorphosed potash deficient rock (hence the andalusite) is responsible for the present assemblage. Pleochroic haloes are abundant in the biotite flakes.

