

DEPARTMENT OF MINES AND ENERGY
SOUTH AUSTRALIA

Rept.Bk.No. 79/129

EXPLANATORY NOTES FOR THE
FROME 1: 250 000 GEOLOGICAL MAP

SHEET SH/54-10 INTERNATIONAL
INDEX.

GEOLOGICAL SURVEY

By

R.A. CALLEN

JANUARY, 1980

D.M. No.: 538/74

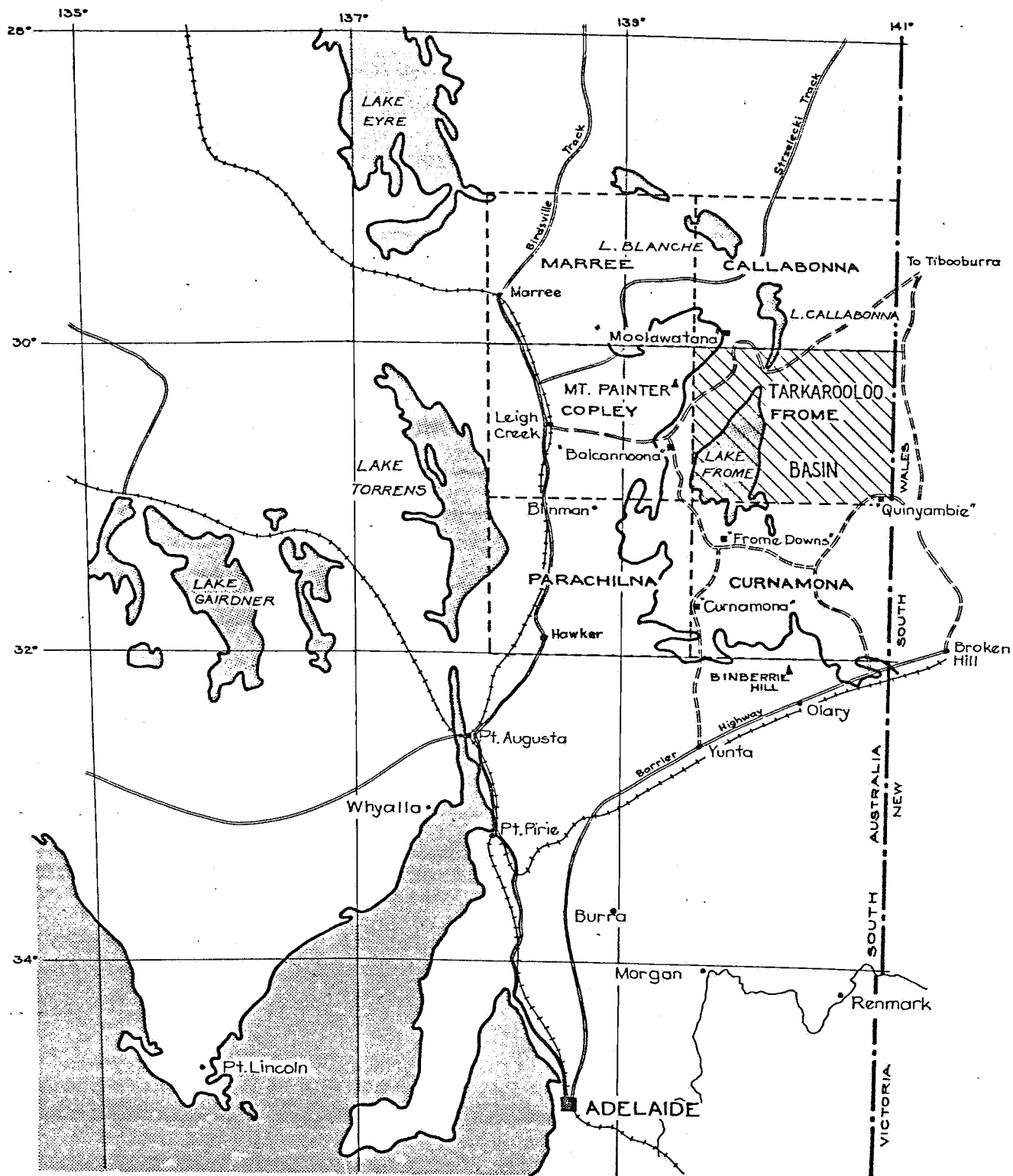
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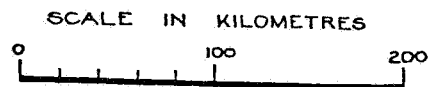
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1:100,000 ENLARGEMENTS

PARALANA	WITTAKILLA	AMERARKOO
PARALANA	COOTABARLOW	COONEE
CALDINA	PUNDALPA	COONEE
FROME		
ARKAROOKA	ELDER	THURLOOKA
FROME	COONARBINE	THURLOOKA
FROME	COONARBINE	QUINYAMBIE

1:63,360 ENLARGEMENTS



Boundary of Cainozoic Sediments (TARKAROOLOO BASIN)

FIG.1

DEPARTMENT OF MINES - SOUTH AUSTRALIA

SCALE: 1:4,000,000

COMPILED: R. Collen

FROME 1:250,000

DATE: 21st April 1976

DRN: A.R. CKD.

LOCALITY PLAN

PLAN NUMBER:

S12171

Compiled by R.A. Callen

INTRODUCTION

The FROME 1: 250 000 geological sheet is located between latitudes 30° S and 31° S and longitudes $139^{\circ}30'$ E and 141° E (Fig. 1) and covers most of the area between the northern Flinders Ranges and the New South Wales border. Its southern and northern limits approximate to the extremities of Lake Frome, which occupies most of the western portion of the sheet. A southern extension of the Strzelecki Desert covers the eastern half of the area.

Main access roads are the Barrier Highway, Port Augusta-Wilpena Pound road (which branches off through Oraparinna, thence to Wertaloona) and the Broken Hill-Tibooburra road (Fig. 1). A track along the New South Wales side of the border fence gives access to the eastern portion of the sheet, but the fence is about 50 to 100 m inside New South Wales, and hence is not shown on the maps. Brougham's Gate is one of the few places where one can pass through the fence into South Australia.

"Putnamutana Outstation", known locally as "Old Moolawatana", and "Quinyambie", are the only permanent habitations, being located in the northwestern and southeastern corners of the sheet respectively. "Frome Downs" is located immediately south of Lake Frome on CURNAMONA. "Balcanoona" and "Wertaloona" are situated along the foothills of the Flinders Ranges to the west of the sheet, on COPLEY. Part of the pastoral lease of "Wirrealpa" is on FROME sheet area.

The climate is dry, hot in summer, with cold winds and occasional frost in winter. Winds are dominantly from the southern quarter with a strong westerly phase during May-August and north winds in summer (Callen, 1974b). Rainfall is

low and erratic (100-125 mm/yr) falling mostly in brief storms with heavier rains of 350-625 mm at 10-20 year intervals. Falls occur mainly during February, March, May, June and December, with April being the driest month. The central portion of the map area east of Lake Frome receives the lowest rainfall, while the highest and most reliable occurs near the Flinders Ranges and in the southeastern corner of the sheet.

The vegetation is sparse. Plant species most useful for geological mapping are *Eucalyptus gillii*, which occurs on the Willawortina Formation (Late Cainozoic) dolomitic sediments adjacent to the Flinders Ranges, particularly where calcrete or carbonates are exposed. *Callitris columellaris* occurs east of Lake Frome and on one island in the lake, tending to prefer the Eurinilla Formation (Quaternary) as a substrate.

The nomenclature used for basins follows Wopfner (1972): the term "Frome Embayment" (Ker 1966, Freeman 1966, Wopfner 1966) refers to the Mesozoic sedimentary basin. "Lake Frome area" refers to the region bounded by the Flinders Ranges to the west, Barrier Ranges to the east and Olary Ranges to the south. The northern limit is taken as an approximate east-west line through the southern portion of Lake Blanche on CALLABONNA. Within the Frome Embayment, the blanketing Cainozoic sediments are unconformable on the Cretaceous and relate to discrete cycle of events, hence the basin of Cainozoic deposition is named the Tarkarooloo Basin (after Lake Tarkarooloo on CURNAMONA). The Poontana Subbasin (after Poontana Ck) occupies the deeper western portion of the Tarkarooloo Basin.

The map was prepared from Department of Lands 60 chain (1: 55,000, Svy. 394,395) black and white photography supplemented by 1: 83,900 colour photography (Svy. 1209).

Landsat band 7 imagery is available (1565-0000-BW7, see Callen 1977). Petrological reports, rock samples and core are on computer file; Core and samples are shared at the SADME Core Laboratories, Frewville. Plans of drillhole distribution (Annual Report of the South Australian Department of Mines and Energy 1973/4, p 26.) and spot height data (plan 7639) are available.

1: 63 360 scale sheets are referred to thus:

Coonarbine, and 1: 250 000 scale sheets thus: FROME.

PREVIOUS WORK

Sturt (1849) was the first to make geological observations on the remarkably parallel sand dunes which he thought were fluviatile deposits. Selwyn (1860) passed through the area in 1859 and referred on his geological map to "recent tertiary deposits on the surface". Brown (1884) made the first geological survey in 1883 of the desert east of Lake Frome, and observed Quaternary conglomerates and clays along the Pasmore River on CURNAMONA. A few years later the Engineer in-Chiefs' Department drilled artesian bores to tap the Great Artesian Basin waters.

The first geological map was published by Woolnough & David (1926). Jack (1930) published cross-sections and defined a number of groundwater basins in the Lake Frome area and described stratigraphy, rock types and basement topography. Kenny (1934) produced a summary of early work and geology in the West Darling district of New South Wales which included a portion of the Frome Embayment.

Osborne (1945) first used the term "Frome Embayment", defined as a synclinal basin bounded by the Flinders and Barrier Ranges. Evans (1946, Evans & Reeves 1948) located ferruginous Tertiary sandstone and Pleistocene outcrops, which he regarded as of Proterozoic age.

Brown (1950, 1953) recorded Cretaceous microfossils.

Ludbrook (in Ker 1966) refined Brown's Cretaceous stratigraphy, and recorded probable Middle Cambrian and Eocene strata. Wopfner (1970) prepared isopach and facies maps of the Cambrian.

The geology and previous work of the Mount Painter block was discussed by Coats (Coats & Blissett 1971) and the portion of this area appearing on FROME was previously published on the Mount Painter Special geological map (Coats *et. al.* 1969). Ker (1966) completed a hydrological survey of the Frome Embayment and prepared a structure contour map of the top of the Palaeozoic strata.

These explanatory notes have been compiled from recent studies on the Frome Embayment and Tarkarooloo Basin by Callen (1974a, 1975, 1976a, 1977) and Callen & Tedford (1976). Other studies relevant to the stratigraphy were made by Firman (1970b).

Geophysical surveys in the Frome Embayment date from a reconnaissance gravity and magnetic survey undertaken on behalf of the Frome-Broken Hill Company Ltd (Kaufman & MacPhail, 1948). This was followed with further magnetic and gravity surveys by Geophysical Services International for Delhi Australian Petroleum Ltd and Santos Ltd (Harding & Geyer, 1963) and by reconnaissance gravity surveys for Delhi-Santos Ltd, conducted by Wongola Geophysical Pty Ltd, during 1964-65. Additional gravity work was completed in 1966 and 1970 (United Geophysical Corporation).

The earliest seismic survey was made by Geoseismic (Aust.) Ltd for Santos Ltd (Dennison, 1960), with more extensive investigations undertaken later by the South Australian Department of Mines (from 1962), United Geophysical Corporation (for Delhi Australian Petroleum Ltd, 1966), and United Geophysical Corporation (for Crusader Oil N.L., 1970 and 1970b).

A summary of much of the above geophysical data is presented in Figure 5, which was compiled from a plan drawn for Crusader Oil N.L. (United Geophysical Corporation, 1970b). The magnetic plan (Fig. 4) is derived from the Frome 1: 250 000 aeromagnetic sheet, an airborne survey by Adastra Hunting Geophysics under contract to the Bureau of Mineral Resources Geology and Geophysics and the South Australian Department of Mines.

Regional interpretations of geophysical data covering the Frome Embayment and adjoining areas have been compiled by Milsom (1965), Westhoff (1968), Parker (1973), and Tucker & Brown (1973).

The potential of the Cretaceous strata for oil and gas in the Great Artesian Basin has been discussed in reports by Wade (1915), Ward (1944) and Osborne (1945). Evans (1946) was the first to suggest that Cambrian rocks of the Frome Embayment might have an oil potential, while Wopfner (1966) and Freeman (1966) assessed the potential of Cambrian and Ordovician strata.

Exploratory drilling for oil in the Frome Embayment began in 1945 with the drilling of Cootabarlow 1, 2 and 3, and Lakeside 1 and Black Oak 1 bores, (the latter on CURNAMONA) by Zinc Corporation Ltd, Enterprise Exploration Ltd and Frome Broken Hill Co. Ltd. This was followed in 1968 by exploration of the Cambrian sediments, when three stratigraphic wells were completed south of Lake Frome on CURNAMONA (Delhi Australia Petroleum Ltd, 1968). Later stratigraphic drilling by the South Australian Department of Mines included the drilling on Frome of Yalkalpo 1 and 2, Mudguard 1, Wertaloona 1, Bumbarlow 1 and Wooltana 1 bores (Callen, 1972; Youngs, 1977, 1978a,b).

Uranium exploration began in 1969 by Kerr McGee Pty Ltd (Ryan 1969), and has continued since.

The South Australian Department of Mines drilled three stratigraphic bores to assist this exploration (Callen 1972, 1976a, 1975). Work has concentrated on the Tarkarooloo Basin Tertiary sediments. A list of open file reports of companies engaged in exploration is appended.

PHYSIOGRAPHY

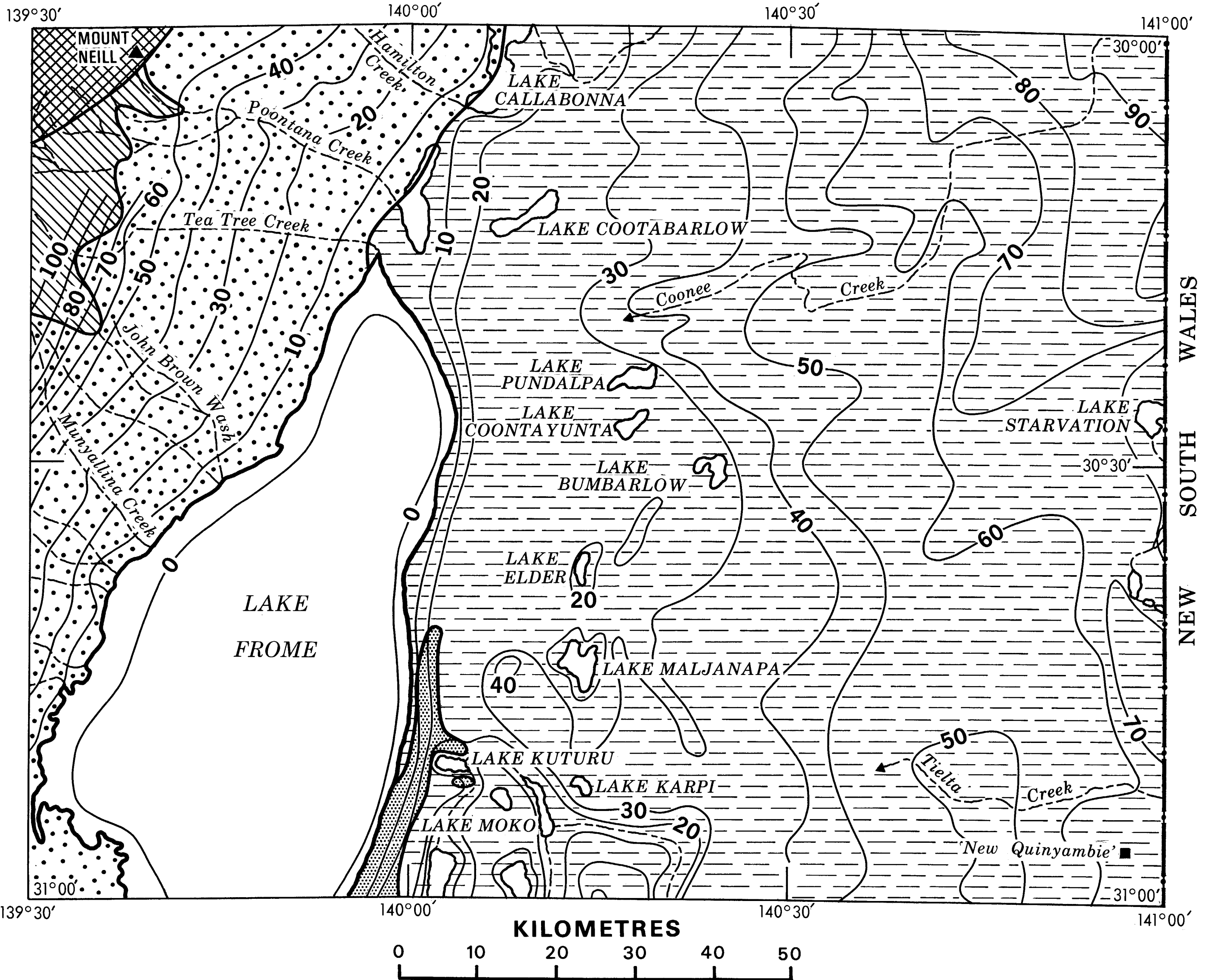
Dominant physiographic regions and topographic contours are shown in Fig. 2.

The Flinders Ranges (region A) in the northwestern corner of the sheet rise abruptly from the Lake Frome Plains (region C) to about 610 m above sea level.

Region B (The Paralana High Plain) rises gently from 100 m above sea level along its faulted eastern edge, to 155 m near the ranges. This plain was uplifted as a unit along with the Flinders Ranges, and originally formed a part of the Lake Frome Plains. The geomorphology is discussed by Twidale (1967).

The Lake Frome Plains (region C) are a complex of coalescing low angle fans, rising from 0.5 m above sea level at the edge of Lake Frome to 70 m near the ranges. They are virtually featureless, apart from occasional creeks with eucalypts, and low gravel ridges. A small area of scattered dunes is developed on *Paralana* and *Caldina* and sparse scrub clothes the area.

Lake Frome (region D) is a playa, probably with some tectonic control. Mound springs are distributed along a straight line near its eastern shore (Draper & Jensen 1976). Islands of stabilised gypsum dunes, belonging to region F, occur in the southern part of the lake. Elevation of the lake bed varies from 0.5 m above to 3.0 m below sea level. Pleistocene shoreline deposits occur sporadically along its western and southern edges. Streams, with watersheds in the Flinders Ranges, contribute detritus to fans along the western shore, where beaches, false spits and bars are developed.



PHYSIOGRAPHIC REGION

A Flinders Ranges

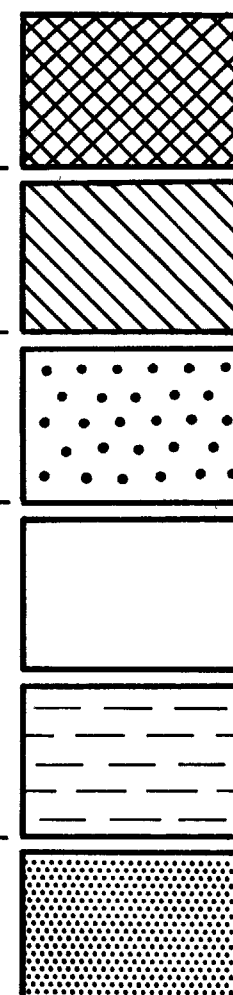
B Paralana High Plain

C Lake Frome Plains

D Lake Frome

E Strzelecki Desert

F Gypsite karst and lunettes



Contour (interval 10 metres)

10

TABLE 1. PRECAMBRIAN AND PALAEOZOIC ROCK UNITS							
Unit	Symbol	Description	Stratigraphic Relationship	Fossils	Thickness (m)	Topographic Expression	Comments
COVER OF CURNAMONA CRATONIC NUCLEUS							
Unnamed Cambrian sediments	7G	Probably Lake Frome Group - reddish brown slate, siltstone, blue grey rubbly limestone (WIRREALPA LIMESTONE) may include Lower Cambrian.		Trilobite		Subsurface, southern Frome Embayment area	Fluviatile and shallow marine
Undifferentiated Cambrian and Proterozoic	G - P	Grey limestone, white quartzite, grey slate				Subsurface Frome Embayment area	

ADELAIDE GEOSYNCLINE & MT. PAINTER INLIER							
Bolla Bollana Formation	Pyb	Massive blue green pebbly greywacke, minor siltstone and quartzite.	Intertongues with Fitton Formation.			Moderate relief	Glacigene
Fitton Formation	Pyf	Tilloid and calc-silicate metasediments quartzites, boulder beds, pebbly arkose.	Unconformably overlies Burra Group.		2000	Low relief	Glaciomarine?
UNCONFORMITY							
Unnamed Dykes		Amphibolite (metamorphosed basaltic rocks)	Intrusive				Basic intrusive
Wywana Formation	Pcy	Actinolitic marbles, minor amphibolite, calcsilicate hornfels and siltstones.	Tectonic contacts			Valleys	Breccia Diapiric?
UNCONFORMITY							
Terrapinna Granite	pEt	Rapakivi - like granite, minor adamellite, augen gneiss.	Intrusive			Low hills	Acid intrusives all related.
Mount Neill Granite Porphyry	pEn	Massive red-weathering porphyritic granite or granite porphyry. Phenocrysts dominant over matrix.	Intrusive			Rounded hills	
Unnamed Granite	pE2	Massive white weathering dark grey microgranite.	Intrusive				
Wattleowie Granite	pEw	Weakly gneissic white granite and adamellite	Intrusive				
Freeling Heights Quartzite (Lower Member)	pEf	Medium to coarse sericitic feldspathic quartzite, quartzose schist, quartz pebble conglomerate at base.	'conformable' contact Brindana Schist		600	Rugged hills, some strike ridges and valleys.	Cross bedding common, rare ripple marks.
Brindana Schist	pEr	Quartzose schist, biotite and muscovite schists, minor garnet sericite schist, epidote quartzite, hornfels, amphibolite, garnet rock.			950+	Forms small ridge	
Pepegoona Porphyry	pEp	Recrystallized porphyritic rhyolite(?)	Contacts may be tectonic		1 300	Low foot hills	Lava flow or near surface intrusive
Mount Adams Quartzite	pEa	Medium grained feldspathic quartzite, minor grit.	Stratigraphic relationships uncertain.		310+	Rugged hills.	
Yagdlin Phyllite	pEg	Laminated grey-green phyllite, lenticular quartzite, (sedimentary lamination).			400		

Note: Plan amended April 1976

FIG.3

DEPARTMENT OF MINES — SOUTH AUSTRALIA			
MOUNT PAINTER BLOCK AND LAKE FROME EMBAYMENT			
PRECAMBRIAN AND PALAEOZOIC			
	R.A.Callen GEOLOGIST	Drn. R.A.C.	SCALE: —
		Tcd. A.F.	74-773
		Ckd.	
Director of Mines		Exd.	DATE: SEPT. 1974

Region F has a unique undulating topography formed of old Pleistocene gypsum lunettes up to 30 m high. It has a gypsum "karst" structure of more recent origin, and lacks vegetation.

There is a prevalence of steep western shores on the lakes and claypans, whereas the eastern shores are shallow and flanked by lunettes.

STRATIGRAPHY

PRECAMBRIAN Table 1

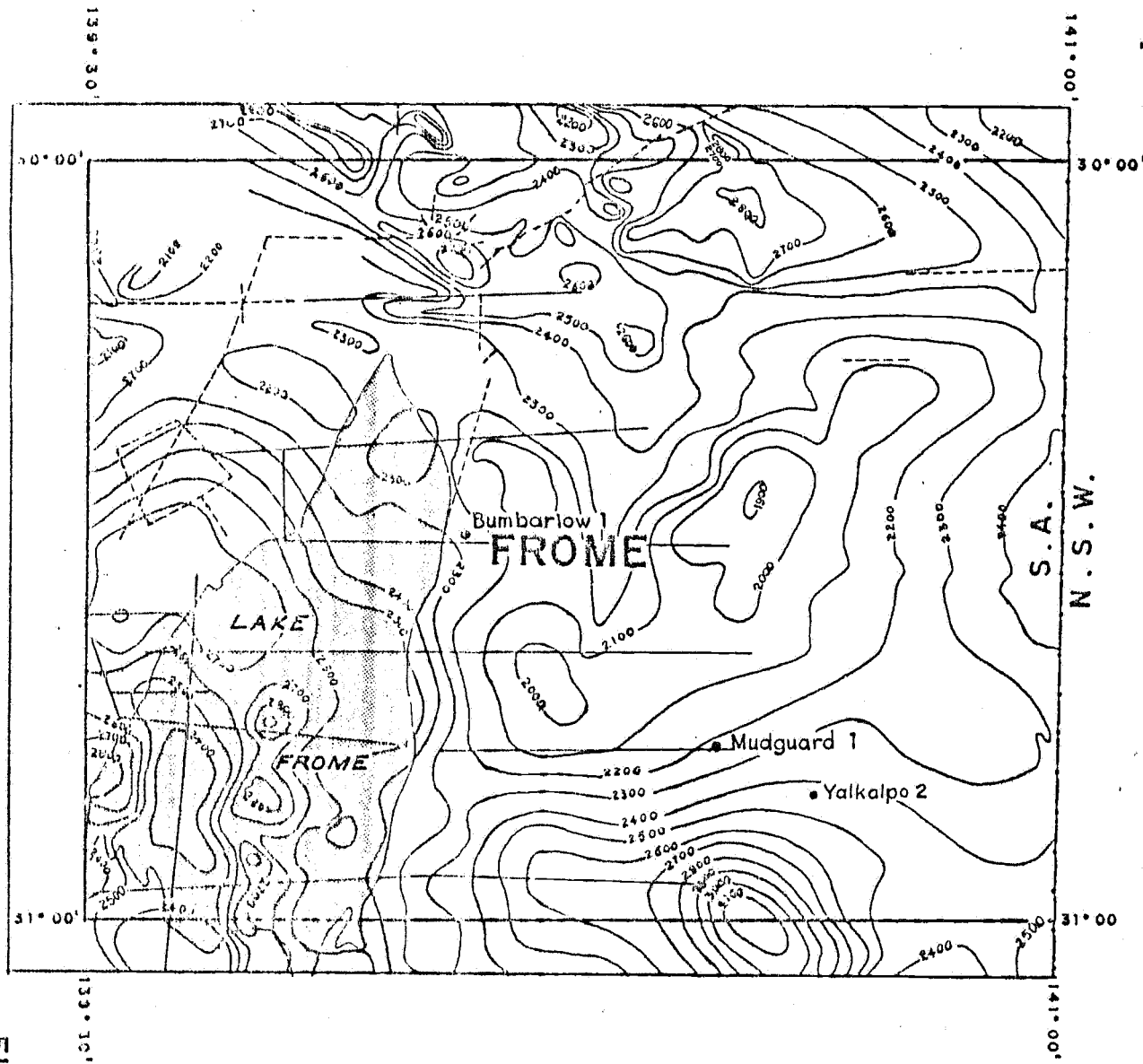
The stratigraphy of the Precambrian is summarised in Table 1 and is derived from Coats and Blissett (1971), the basement geology on FROME being adapted from the Mount Painter Special geological sheet (Coats *et. al.* 1969). A list of unpublished University theses is given by Blight (1977), and geochronology is discussed by Webb *et. al.* (1978).

Within the Terrapinna and Wattleowie Granites, overprints have been used on FROME to differentiate bands which are considered by Coats to represent metasediments. Young (1973) discusses these rocks ("microgranites") and the contacts of the Pepagoona Porphyry.

A few small outcrops of the "Younger Granite Suite", included in the area of brecciated Wywyana Formation, are not differentiated on FROME.

In the subsurface, little information is available. Crystalline basement rocks of the Curnamona Cratonic Nucleus (Thomson 1974, Thomson *et. al.* 1976) probably occupy most of the map area and are most likely continuous with the Mount Painter Block (see also Sprigg's "Paralania", Sprigg *et al.*, 1958). An extensive porphyry body occupies at least the southern part of the Benagerie Ridge on FROME and is best represented in Mudgard 1 bore (Youngs 1977). The porphyry is not reflected by features in the magnetic or gravity maps (Figs. 3, 4) but is represented by

FIG.3



R. A. CALLEN — Geologist

TOTAL MAGNETIC INTENSITY

(From FROME 1:250 000 aeromagnetic map and selected seismic lines).

Magnetic intensity at 100γ intervals... 2100

CRUSADER OIL N.L.

S.A. Dept. Mines

S.A. Dept. of Mines and Energy
bores drilled since publication
of the map.

Scale in Kilometres

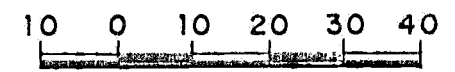
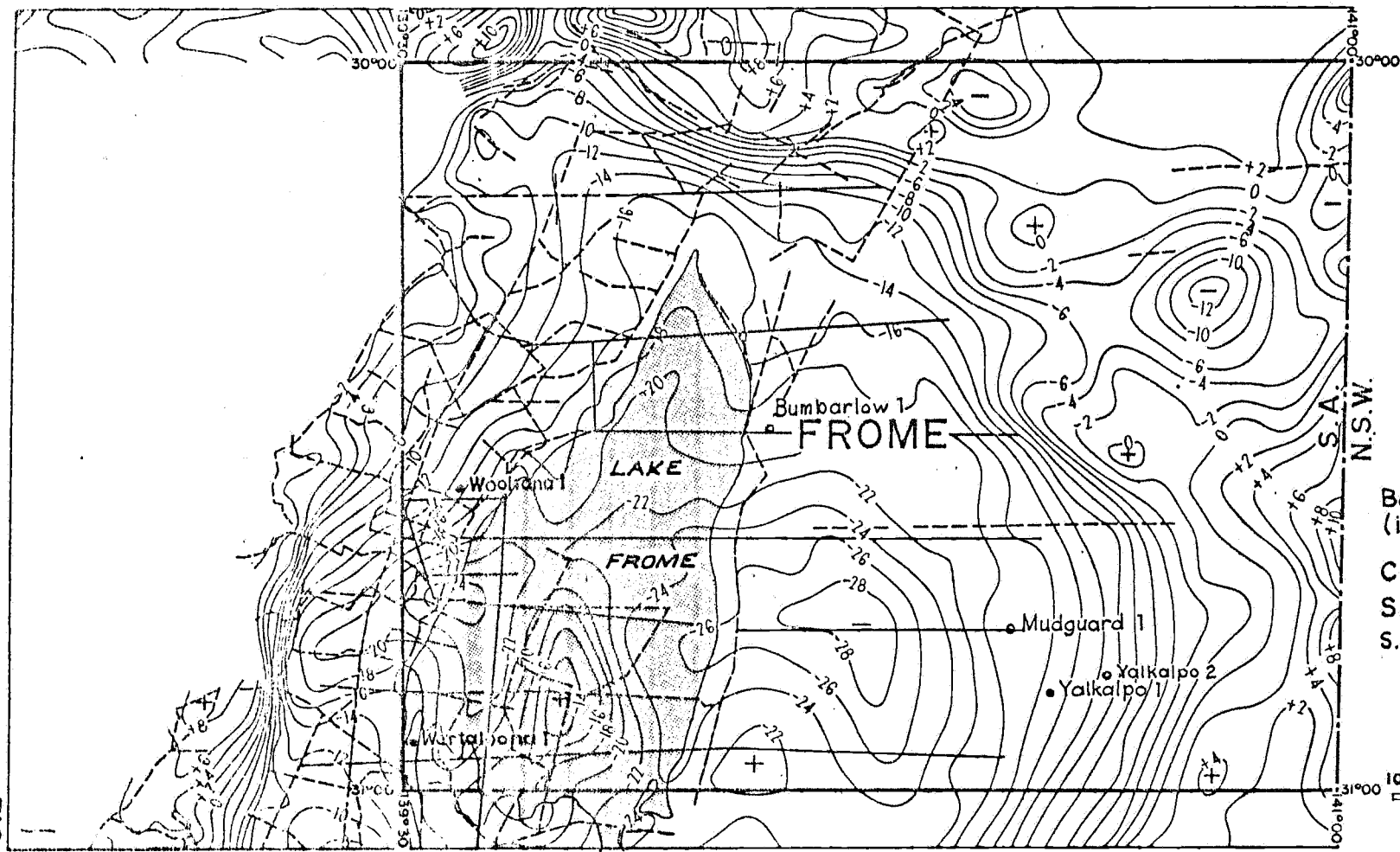


FIG.3
S12011

S.A. DEPT. of MINES

FIG. 4



BOUGUER GRAVITY AND SEISMIC LINES

(Compiled from plan drawn by
United Geophysical Corporation
for Crusader Oil N.L., open file
envelope 1275(v) Dept. of Mines
and Energy.)

- Bouguer gravity contour
(interval 2 milligals).....
- Crusader Oil N.L.
- S.A. Dept. of Mines
- S.A. Dept. of Mines and Energy bores
Drilled prior to publication of map.....
- Drilled since publication of map

SCALE IN KILOMETRES

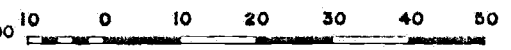


FIG. 4
S 12010

S.A. Dept. of Mines

R. A. Callen — Geologist

a characteristic lack of reflections in seismic profiles. Thomson (1966, p. 216) noted similarities to the Gawler Range Volcanics (Blissett 1975), but ages are younger (1160-1500 Ma) and similar to basaltic volcanics in Bumbarlow 1 bore (Youngs 1978b).

Depths to magnetic basement, interpreted as crystalline basement by Tucker and Brown (1973), permit a thin Adelaidean sequence of 500-600 m over the southwestern portion of FROME. Bumbarlow 1 bore intersected 300 m of Precambrian basaltic lava flows, interbedded with red and green shale, sandstone and conglomerate. Clasts include metamorphic and volcanic rocks. Below this 18 m of grey laminated shales. Youngs (1978b) suspects the contact with the shales is unconformable.

Unmetamorphosed purple cross-bedded coarse sandstones, cut by thin quartz veins, are located on the New South Wales border fence at latitude $31^{\circ}40'$ (Callen, 1974b). These are probably Adelaidean.

Thus a thin relatively undeformed platform cover of Proterozoic, (?)Adelaidean rocks exists on FROME. PALAEOZOIC (Table 1).

There is no exposure of Palaeozoic rocks, records being limited to two stratigraphic bores, Yalkalpo 1 (Callen 1972) and 2, (Youngs 1977, 1978a,b), east of the Benagerie Ridge. Yalkalpo 1 bore intersected a few metres of Lower Cambrian (personal communication, B. Daily, University of Adelaide, 1975) grey limestone above the Precambrian porphyry, whereas Yalkalpo 2 bore, intersected 540 m of Lower and Middle Cambrian red-beds and blue-grey limestone, without reaching basement. The absence of Cambrian in Bumbarlow 1 bore suggests the Arrowie Basin (Wopfner 1970a,b) is probably restricted by the Benagerie Ridge to the west of Lake Frome. The Cambrian sequence east of the Benagerie Ridge can probably be regarded as a separate basin.

COMPILED: R. A. Collen		DEPARTMENT OF MINES—SOUTH AUSTRALIA		SCALE: (Chart)	
DRN:		CKD.		DATE: 5 TH APRIL 1976	
LAKE FROME AREA				PLAN NUMBER	
MESOZOIC ROCK UNITS				S12170	

TABLE 2 LAKE FROME AREA - MESOZOIC ROCK UNITS

ROCK UNIT	SYMBOL	LITHOLOGY	FOSSILS	THICKNESS (METRES)	COMMENTS
MARREE SUB-GROUP (MARREE FORMATION) (CODNADATTA FORMATION AND BULLDOG SHALE equivalent)	K1r	Monotonous sequence of dull greenish grey micaceous shale and siltstone, intraformational breccia, minor pebble beds.	Burrows, leaf impressions & carbonaceous material. Spores. Foraminifera	150-275	Often pyritic or ferruginous. Present in most deep bores, outcropping in small areas along Flinders Ranges. Sand beds may be Attraction Hill Sandstone in part. Could include some winton Formation at top (non-marine). The Codnadatta & Bulldog shale are lithologically indistinguishable or present evidence.
CADNADATTA FORMATION, formerly (Pelican Well formation)	K1c	Sandstone, micaceous, medium grained and subangular; interbedded with dull greenish-grey micaceous shale and silty shale; pebble and boulder beds common, some limestones.	Leaf and stem detritus, spores.	50 max.	Identified in Cootabariu No. 2, Coonee Ck., Arcoola Lakeside No. 1, Black Oak, and Curraworra Bore and Yalkalpc No. 1. Basal part Jurassic and partly equivalent to the Algebuckina Sandstone.
UNNAMED CONGLOMERATE	K1l	Marine boulder bed of granitic rock types, rhyolite porphyry and grey slate. Low percentage grey slate matrix.		2-4 min.	Probably Jurassic, derived from Benagerie Ridge. Porphyry locally derived.

CRETACEOUS

JURASSIC

JURASSIC

Delhi-Santos Lake Frome 1, 2 and 3 bores were drilled on CURNAMONA along the southern edge of Lake Frome. They intersected 780 m of Middle Cambrian strata, and by comparison with exposed sections on PARACHILNA, almost penetrated the base of this sequence. Other bores on CURNAMONA also intersected Cambrian rocks (see open file envelopes listed in the appendix, for the western part of the sheet area). These thick Cambrian sections west of the Benagerie Ridge on CURNAMONA presumably extend onto FROME. The "base of Mesozoic" on the tectonic sketch, was determined from seismic data, and does not fully accord with more recent private company drilling west of Lake Frome.

A gravity low in the southeastern portion of FROME suggests some low density sediments, possibly Cambrian, are present. Wopfner (1966), showed that Ordovician strata might underlie the eastern margin of FROME. Later Palaeozoic sediments are unknown.

MESOZOIC (Table 2).

Terminology for the Mesozoic of the Frome Embayment was first introduced by Forbes (1966). The following information is from subsurface data, there being no outcrop.

The oldest Mesozoic unit on FROME is the unnamed ?Jurassic conglomerate in Yalkalpo 1 bore, resting on the Lower Cambrian ?Parara Limestone. This has a very small proportion of grey silty shale matrix, identical to the overlying Cadna-owie Formation equivalent. Clasts are porphyry, granite and various metamorphic rocks derived from the Benagerie Ridge (Section A-A' on map).

The Cadna-owie Formation (Wopfner 1969, Wopfner *et. al.* 1970, now used in preference to Pelican Well Formation in the Frome Embayment), is gradational with the Marree Formation. However, in Yalkalpo 1 bore it can be identified using the petrophysical logs.

The Cadna-owie Formation on FROME is shale, sand, and silt with lenses of coarse clasts ranging in size from pebbles to large boulders. Sandy beds are common, especially near the base. Clasts include white quartzite and sandstone, reddish brown porphyry, and fine siliceous shale. Some of these may have been reworked from Permian sediments, which have mostly been stripped. The work of R. Morgan (personal communication 1974, Geological Survey of New South Wales) indicated the basal Cadna-owie Formation was deposited during the Late Jurassic.

The Marree Formation (Bumbarlow 1, Cootabarlow 1 bores) constitutes a monotonous sequence of dull greenish grey to bluish-grey finely micaceous shale and siltstone, laminated beds, intraformational breccia and minor polymictic conglomerate. Much plant leaf and stem material and numerous invertebrate burrows or traces are present. Bivalves, foraminifera and microflora are common. No weathering horizon is developed on the Cretaceous in subsurface. Morgan deduced from palynological evidence that the upper part of the Marree Formation had been eroded in Yalkalpo 1 bore. Unlike some other areas (e.g. Oodnadatta region) the Marree Formation in the Frome Embayment cannot be subdivided at present.

The Cretaceous Winton Formation has not been positively identified on FROME. Ludbrook (in Ker, 1966) shows the Cenomanian is absent from the sequence south of Tilcha Bore on CALLABONNA. However, thin Winton Formation may occur in the extreme north of the sheet area (Muloowurtina and Poontana bores). A brownish sandy sequence in Cootabarlow 1 bore and sandy beds at the top of Poontana bore may represent the unit, but have not been distinguished as such on the bore logs or sections.

Spores and microplankton have been studied by Dettman (1963), Dettman & Playford (1969, especially pages 182-192, tables 9.1,

TABLE 5 LAKE FROME AREA - CAINOZOIC - TERTIARY

AGE	ROCK UNIT	SYMBOL	LITHOLOGY	SEDIMENTARY STRUCTURES	CRITERIA FOR CORRELATION	FOSSILS	GEOMETRY	GEOMORPHIC EXPRESSION	SECONDARY WEATHERING EFFECTS
Uncertain LATE MIOCENE-PLIOCENE possibly to EARLY PLEISTOCENE From relation- ship to NAMBA FORMATION	WILLAWORTINA FORMATION	Czw	Extremely to very poorly sorted bouldery to pebbly silt clay or sand clay with numerous greenish carbonate nodules enveloping matrix and skeleton grains in southern areas. Lenses of pebbles to large boulders common near Flinders Ranges, sand beds further east. 2.5-7YR5-8/1-8, 5-7Y6-8/1-8.	Crude wavy and lenticular horizontal lamination. Well defined medium scale cross-beds. Beds 0.5-1 m thick. Fining upwards cycles in lower part.	Greenish clays with very poorly sorted sand to boulder content, usually matrix supported framework. Greenish silty carbonates. Red brown mottling common.	Rare vertebrates	Wedge shaped mass, thinning east, between Lake Frome and Flinders Ranges, and to south. (0-150m) Thin sheets	Forms basic landscape of uplifted "high plain" flanking Flinders Ranges, where it supports growth of <i>Eucalyptus gillii</i> . Crops out as vertical or overhanging cliffs along creeks. Cliffs along Pasmore River and south shore of Lake Frome.	Well developed massive white hard carbonate developed in long cylindrical or horizontal sheets, pisolitic structures. Hard buff groundwater calcrete cements porous sediment. Massive crusts of gypsum rosettes. Red-green mottling, green colour increasing with depth, black patch stain on fractures.
		CZst	Brown silcrete and iron oxide cement		Colour, petrology			Capping on sands of Tmb and Tee	
Medial from plant spores and pollen, to ? Late Miocene.	NAMBA FORMATION	Tmb	Alternating fine to medium poorly sorted sands, silt and clay, thin dolomite and limy dolomite beds, often oolitic. Sands generally angular, of high sphericity, but may also be polished and subrounded with low sphericity. Clays may be black (ferrous iron) and tough with characteristic irregular shiny-surfaced fractures, (unit Tmb ₂) or pale green or grey. Sands and clays may be carbonaceous. Dolomites often have nodular structure, and are penetrated by branching pores lined with manganese. Rare charales limestone. 1-10Y2-9/0.25-6, X0-N9, 10YR2-7/1-2, rarely 5B5-7/1, 5-10R5-3.5/3-4, 5GY4-7/0.5-1 and 5G6/1.	Horizontal lamination common, especially in silt and very fine sand. Very small to medium scale cross-bedding common in silts and sands. Irregular oolites in carbonates, often rod shaped. Intraformational brecciation, quick flow structures common. Bioturbation common. Rare shrinkage (2 sub-aqueous) cracks and slumps. Cyclic deposition in unit Tmb ₁ .	Pale grey green clays and fine silts and sands, dolomites, tough black clay with irregular fractures. Fish spines in lower beds. Vertebrates - lung fish teeth.	Ostracoda (Cypridids), Charales oogonia and stem moulds, <u>Potamopyrgus</u> s.l., Aves, Cetacea Dipnoi, Teleostei, Chelonii Crocodilia Monotremata Marsupialia Plant spores - <u>Nothofagus</u> , <u>Dacrydium</u> , <u>Podocarpus</u> , <u>Eucalyptus</u> , GRAMINEAE RESTIONACEAE CUPANIEAE <u>Pediastrum</u>	Thin widespread blanket and shallow basins (20-250 m)	Sloping cliffs, often capped by dolomite, gypcrete or secondary carbonate cemented younger units, along edges of dry lakes and creeks, east and south-east of Lake Frome. Less commonly as dipping sequences upfaulted along western margin of Flinders Ranges. Mainly subsurface.	Sometimes capped by ferruginous yellow-brown orthoquartzite silcrete. (CZst) Alunite soil-like horizons associated with uppermost black clay. Massive gypsum crust. Reddish mottling (weakly developed) Also capped by puddingstone opaline and chalcedonic silcrete.
PALEOCENE AND EOCENE from plant spores + pollen	EYRE FORMATION	Tee	Moderate to poorly sorted sands, grain size ranging from fine to small pebbles, often polished, subrounded and of low sphericity. Polish greater for larger sizes. Characteristic basal pebble beds with milky, clear or smoky quartz, grey quartzite, black chert, red jasper, agate, may be very micaceous. Brown to grey colour common from presence of carbonaceous matter. Dark brown to black carbonaceous clay, silty, soft. EOCENE beds tend to be silty and carbonaceous, PALEOCENE coarser grained. 5YR2/1, 10YR6/2, N1-N8, 4Y6/0.45, 5B8/1, 5B6-7/2.	Medium scale and small scale trough cross-stratification and planar cross-stratification. Lamination. Tool marks. Fining upwards sequences.	Polished basal pebble bed. Essentially sands, often carbonaceous, black clays are soft and silty with much carbonaceous matter.	Spores pollen, dinoflagellates, arthropods (rare), leaf, stem and fruit moulds. Plants include fragments carbonized wood. <u>Podocarpus</u> <u>Dacrydium</u> <u>Nothofagus</u> PROTEACEAE CASUARINACEAE GRAMINEAE RESTIONACEAE CUPANIEAE MALVACEAE SANTALACEAE MYRTACEAE HALORAGIS	Thin widespread blankets (10-75m)	Silicified outcrop upfaulted on margins of basin. Isolated outcrop on state border. Usually moderately to steeply dipping.	Often ferruginized and capped by massive microcrystalline quartz silcrete with botryoidal or columnar structure. (Tsi & CZst) Frequently silicifies the sediment itself.

After H. Wopfner et. al. (1974), Callen & Tedford (in press)

FIG.7

Note: Plan amended April 1976

DEPARTMENT OF MINES — SOUTH AUSTRALIA			
LAKE FROME AREA — CAINOZOIC TERTIARY			
	R.A. Callen GEOLOGIST	Drn. R.A.C. Tcd. A.F. Ckd. Exd.	SCALE: — 74-769 DATE: SEPT. 1974
Director of Mines			

9.3, 9.4), Cookson & Eisenack (1958) and Eisenack & Cookson (1960), Evans (1966). Cootabarlow 2 bore is the reference section for the *Crybelosporites stylosus* zone of Dettman.

(N.B. On the bore log for LC1A (in northwestern part of map area), the thickness for the Cretaceous (K) should read 120 m. The basal conglomerate and lower part of the Cadna-owie Formation are Jurassic not Cretaceous as shown in the map legend).

CAINOZOIC

Several new names have been used for rock units in the eastern side of the Flinders Ranges, north of the Olary Block. These units are defined elsewhere (Wopfner *et. al.* 1974, Callen & Tedford 1976). Suggested equivalents are given in Table 5. Tertiary (Table 3 and section D-D' on map).

The new name TARKAROOLOO BASIN is given to the Cainozoic sediments deposited in the Lake Frome area. The northern limits of the basin are at present undefined. These continental deposits form a thinner and more extensive cover than the Frome Embayment Mesozoic sequence. Sedimentation was initiated by a different series of tectonic events than for the Mesozoic rocks and the locus of deposition is nearer the Flinders Ranges. There are considerable differences in the tectonic history of the Palaeogene strata here compared with those of the Neogene.

Palaeocene-Eocene

Deposition during these epochs is represented by the EYRE FORMATION (Wopfner *et. al.* 1974, formerly Murnpeowie Formation of Forbes 1966), recognised over a wide area of the Great Artesian Basin, and well-developed in the Lake Frome area. Supplementary reference sections for the Eyre Formation in the Tarkarooloo Basin are Mines Administration LC1A bore on *Paralana* (FROME) and the Reedy Springs section on MARREE (Wopfner *et. al.*

1974). Yalkalpo 1 bore is a third important section, representative of the reduced sequence on the Benagerie Ridge.

The lithology of the Eyre Formation is consistent over the whole FROME 1: 250 000 sheet, with fine to very coarse-grained (medium dominating) sub-mature to mature sands, moderately to poorly sorted and frequently bimodal. The sands are often pyritic and carbonaceous, becoming micaceous, kaolinitic and angular towards the Olary Ranges. Quartz grains are subrounded and original crystal faces often visible. Matrix is usually dominated by kaolinite, with accessory montmorillonite and illite. A characteristic pebble bed is present at the base, with black chert (both fibrous chalcedony and micro-crystalline quartz), fossil wood, red 'jasper', and quartz of milky, grey, yellow or no colour. The grains are characterised by high polish. Lenses of black carbonaceous silt occur.

The Eyre Formation reaches its maximum thickness in the Poontana Sub-basin, thinning over the "high" in the eastern part of the sheet. The base of the Tertiary was contoured for the tectonic sketch using sparse bore data. A structure contour map using a selection of extensive company drilling is included in plan 76/6.

In the Lake Frome area there is excellent biostratigraphic control (Harris *in* Wopfner *et. al.* 1974), discussed in detail for the Mooloowatana-Reedy Springs area by Callen (1975, Fig. 15). A disconformity may occur within the unit, the late Palaeocene being unrepresented. The Eocene is typically finer grained and of a characteristic dark brown colour, with leaf remains, in contrast to the coarse sandy Palaeocene deposits with black carbonaceous matter.

Conglomerate which crops out on the New South Wales border may be the only exposure on FROME (*Thurlooka*). It represents basal Eyre Formation.

In subsurface, the contact with the Cretaceous strata is well defined: coarse polished quartz sand with chert pebbles overlies dull green grey silty shale of the Marree Formation. The upper contact may be difficult to identify because sands may occur in the overlying Namba Formation (Callen & Tedford 1976 a,c, Callen 1976a) as in Yalkalpo 1 bore. In general, the first sands encountered below the black clays or carbonate horizons of the Namba Formation are those of the Eyre Formation. Eyre Formation sands are usually coarser grained, more highly polished, with higher roundness and lower sphericity than those of the Namba Formation.

Both the lower and upper contacts show regional disconformable relationships. Channelling is poorly developed or absent at the lower contact on FROME (see Wopfner *et. al.* 1974), though channel sands of the Namba Formation may cut into the Eyre Formation.

Deposition of the Eyre Formation ceased during the Upper Eocene, leaving a widespread sand blanket.

Oligocene - Early Pliocene

Silcrete and bleaching of ?Oligocene age, occur to a limited degree on dipping sequences around the margins of the Tarkarooloo Basin, although they do not occur on FROME itself. Incipient silicification and some silcrete ?pebbles are present in Yalkalpo 1 bore.

The Namba Formation (Callen & Tedford 1976) was deposited essentially during the Early to Medial Miocene although it may extend into the Late Miocene. The name is derived from Lake

Namba (new name*) on CURNAMONA. The type section is in Yalkalpo 1 bore on *Quinyambie*, and an important supplementary section is located at Wooltana 1 bore on *Arkaroola*. The outcrop supplementary section is in a cliff on the central western shore of Lake Tarkarooloo* (CURNAMONA). The unit is the time equivalent of the Etadunna Formation (Stirton *et. al.* 1961), as shown by similarity in certain characteristic fossil-vertebrates, and comparable microflora (Harris 1972).

A Medial Miocene age (?Balcombian-Batesfordian) was determined palynologically (Harris, personal communications 1972, 1973), the flora being comparable with that of the Munno Para Clay Member (Lindsay & Shepherd 1966, Lindsay 1969) of the St. Vincent Basin. The Miocene microflora was found in Cootabarlow 2, Wooltana 1, LC1A, F22/26, and Glenmore 1 (on BROKEN HILL, see Brunner 1967) bores, and in Lake Eyre 20 bore in the Lake Eyre Basin. The Lake Eyre bore microflora is very similar to that of Wooltana 1 bore. Correlation with foraminiferal zones of the Murray Basin and southern Victoria has been obtained through Tricentrol Aust. Ltd bore near Mutooroo, OLARY (Lindsay & Harris 1973). Vertebrate fossils of considerable evolutionary and palaeogeographical significance are described by Tedford *et. al.* (1977). Although the possible age range of the Namba Formation is Late Oligocene to Late Miocene the unit was probably deposited in a relatively short time interval in the Early to Medial Miocene.

The Namba Formation is divided into two units on the map, though in Wooltana 1 bore it may be divided into six (Callen & Tedford 1976) using geophysical and litholog-data. The regional two-fold subdivision is based on the occurrence of dark grey, tough clays with a characteristic irregular shiny fracture

*Names formalised by the Geographical Names Board, South Australia, 1973. Detailed list in Callen (1974b).

pattern in the lower unit (Tmb₂). This unit is dominated by clays rich in smectite, and typified by rather poorly defined cyclic deposition involving black clay, cross-bedded sand, finely laminated silt and thin carbonates. The upper unit (Tmb₁), also containing carbonates, is dominated by fine laminated silt and cross-bedded fine sand with prominent bioturbation. Muds are dominated by illite and are light green grey or olive. The numbering for the two members (Tmb₁, Tmb₂) has been revised to the correct sequence in Callen & Tedford (1976), Tmb₁ being the older member and Tmb₂ the younger.

The carbonates (dolomite) alternate regularly with clays, and are rich in palygorskite at the base of each major carbonate horizon. Clay mineralogy was determined by Brown (see list of Brown's work in Callen 1975, Appendix 2), and is discussed in Callen & Tedford (1976) and Callen, (1977).

A distinctive carbonaceous laminated dark green to black fissile claystone is restricted to the Poontana Sub-basin (Wooltana 1 bore Central Pacific Minerals Pty Ltd bores see Schindlmayr 1970). It contains gastropods (M. Buonaiuto pers. comm. 1978), fish remains, ostracods and microflora.

The sediments reach their maximum thickness in the Poontana Sub-basin, and appear to retain the same thickness up to the marginal fault system of the Flinders Ranges. There is little information on the Tertiary close to the ranges, west of the Poontana structure. Holes drilled for Exoil and Western Nuclear (Aust.) Ltd suggest unit Tmb₁ is missing (cross-section A-A') in the outer rim of the Paralana High Plains (region B, Fig. 2).

The outcrop of FROME is mainly of unit Tmb₂, black clay or silicified sands. Most outcrop occurs on *Eurinilla* and

Siccus on the northern edge of CURNAMONA, where unit Tmb₁ is also present and a sharp undulating contact with Tmb₂ can be locally observed.

An alunite horizon, developed in the upper part of unit Tmb₂ is widespread in the western high plains on COPLEY and east of Lake Frome at Lakes Starvation and Bumbarlow. Its structure suggests it is related to a weathering horizon or soil profile, thus indicating an hiatus.

Late Tertiary to Early Quaternary

Late Miocene - Early Pleistocene.

There is a ferruginous horizon resembling the Karoonda Surface (Firman 1971) developed on the Namba Formation. This is included with unit CzSi on the map, and is probably equivalent to the ferruginous and silicified paleosol complex of ?Pliocene age recognised by Firman (pers. comm. 1979).

Upon the Namba Formation clays is developed a siliceous horizon, consisting of microcrystalline quartz, with numerous cavities of pedogenic origin infilled by opal and chalcedony. On sandstones of the same unit this is represented by quartz overgrowths on the detrital grains or microcrystalline quartz cement, producing massive "grey billy". This silicification came after ferruginisation, hence the silcrete duricrust generally has a red and white mottled appearance. Its frequent association with alunite suggests it is older than or equivalent to the Willawortina Formation.

The Willawortina Formation constitutes a wedge shaped deposit, thickest near the Flinders Ranges on COPLEY, of poorly sorted conglomerates, nodular to massive carbonates, and red and green mottled clays, interpreted as flood plain and stream deposits. On FROME it is restricted to the Poontana Sub-basin. Its type section is WC2 bore on the Paralan High Plains with an

outcrop supplementary section, south of "Wertaloona" on *Balcanoona* (COPLEY). Another supplementary section is Wooltana 1 bore. The Willawortina Formation intertongues with the upper member of the Namba Formation (Tmb_1). The base of the unit is well defined by petrophysical logs in the southern portion of the basin, but in the northern area, where secondary carbonate cementation is absent, it is less distinctive. The boundary with the older Tertiary sequence has been mapped by Leeson (1967, see also Coats *et. al.* 1973) on *Balcanoona* (COPLEY). A thick sequence of conglomerate and clay of the Willawortina Formation overlies clastics of the Namba Formation with no apparent evidence for an unconformity in the Beverley area (Paralana High Plain).

The Willawortina Formation was referred to the Avondale Clay (Firman 1967) on the adjoining COPLEY sheet.

In the area flanking the northwestern shore of Lake Frome, drilling has intersected hard red-brown calcareous and green-grey clays (Schindlmayr, 1970). These are probably fine grained equivalents of the Willawortina Formation which grade into the Namba Formation unit Tmb_1 .

The Millyera Beds (Qp_6) comprise laminated light green to bluish green clay with thin fine sand beds and thin charophyte-stem limestone and gypsum laminae. They are locally disconformable on the Namba Formation. Abundant ostracods, charophytes, gastropods, bivalves (M. Buonaiuto, personal communication, 1978), and worm burrows are present. The Beds were named after Lake Millyera on *Siccus* (CURNAMONA), where they are best exposed. When unweathered they are finely laminated and contain abundant ostracods, and lenses of sand, rich in charophyte oogonia. The clays have a characteristic sub-conchoidal fracture, distinguishing them from the Miocene sequence, and often have a distinctive finely laminated charophyte limestone or gypsum bed at the top.

In Lake Callabonna, clays similar to those of the Millyera Beds contain vertebrate remains of Plio-Pleistocene age: these are the "*Diprotodon*-bearing beds" of Tedford (1973).

Unit 'Qp₅' crops out in two widely separated areas: along the northeastern shore of Lake Millyera and on the northern shores at the junction of Lake Frome with Lake Callabonna. Unit Qp₅ invariably occurs directly overlying the green clays of the Millyera Beds, and its type area is on the southeastern shore of Lake Millyera. Lithologically it is fine to coarse fossiliferous sand. Fine cross-bedding and wavy horizontal bedding is typical, and abundant charophyte stem fragments and oogonia occur. Bright reddish orange gypsum-impregnated silt lenses are present, closely resembling the Eurinilla Formation (described below).

The Millyera Formation of Callen & Tedford (1976) constitutes the Millyera Beds (Qp₆) and unit Qp₅ combined, though recent work suggests the original practise of separating them is probably correct, and a disconformity may exist between.

The next youngest unit is the unnamed conglomerate Qp₇, consisting of extensive sheets of planar cross-bedded conglomerate and sand, exposed along the southern shore of Lake Frome and the Pasmore River. The unit is solidly cemented with secondary white or buff carbonate. It contains boulders and pebbles derived from the Olary and southern Flinders Ranges, with a local admixture of carbonate nodules derived from the Willawortina Formation. The Willawortina Formation contains numerous similar lenses of conglomerate, also cemented, which may be difficult to distinguish from unit Qp₇ in isolated outcrops. The unnamed conglomerate may be equivalent to unit Qp₅, by comparison with a similar sequence in Lake Tarkarooloo on CURNAMONA, which grades into unit Qp₅ northwards along the outcrops on the lake floor.

TABLE 5

CAINOZOIC UNITS OF FROME AND POSSIBLE EQUIVALENTS

TIME UNIT	TARKAROOLOO BASIN AFTER CALLEN & TEDFORD (1977)		EQUIVALENTS IN OTHER BASINS AND ADJACENT MAP SHEETS		
	SYMBOL	NAME	FIRMAN 1971 "FROME EMBAYMENT" AREA	COATS et. al., (1973) COPLEY	STIRTON et. al., 1961 LAKE EYRE
HOLOCENE (RECENT)	Qra		Qra	Qra	
	Qri	Unnamed lake deposits	Qri	Qri	
	Qrs	SIMPSON SAND	SIMPSON SAND	Qrs	
	Qrb	Unnamed clays & sands	TINGANA CLAY CALLABONNA CLAY LOVEDAY SOIL	ARROWIE FORMATION	Qri
	Qcc1	Unnamed soils			
LATE PLEISTOCENE TO HOLOCENE	Qp1	COONARBINE FORMATION + Older Holocene fan deposits	POORAKA FORMATION	POORAKA FORMATION	
	Qp2	Unnamed gypsum lunettes			
	Qcc2	Unnamed soil	LOVEDAY SOIL		
	Qp3	Unnamed beach deposits			
	Qcs	Gypsum in Eurinilla Formation	Qcs	Qcs	
MEDIAL TO LATE PLEISTOCENE	Qp4	EURINILLA FORMATION	POORAKA FORMATION + TELFORD GRAVEL	POORAKA FORMATION	KATIPIRI SAND & TIRARI FORMATION
	Qca	Unnamed calcrete	BAKARA SOIL + (RIPON CALCRETE) Includes groundwater calcrete in basal Eurinilla Formation.	Qca	
	Qp6 + Qp5	MILLYERA FORMATION (MILLYERA BEOS + unit Qp5)			
	Qp7	Unnamed Conglomerate (Equiv. MILLYERA FORMATION)	TELFORD GRAVEL	TELFORD GRAVEL	
	CZw	WILLAWORTINA FORMATION	AVONDALE CLAY & "Conglomerate at Lyndhurst"	BUNGUNNIA LIMESTONE Equiv. "T-Q" CLAY,	
LATE MIOCENE TO EARLY PLEISTOCENE	CZsi	Ferruginous horizon & Silcrete (Relationship to Willawortina Formation, uncertain)		Not recognised	
	Tmb	NAMBA FORMATION	AVONDALE CLAY	BUNGUNNIA LIMESTONE AVONDALE CLAY	WIPAJIRI FORMATION ETADUNNA FORMATION
MEDIAL TO LATE MIOCENE	Tsi	Not represented	Tsi	Tsi	
OLIGOCENE ?					
EARLY PALEOCENE TO LATE EOCENE	Tee	EYRE FORMATION	EYRE FORMATION	EYRE FORMATION	

DEPARTMENT OF MINES - SOUTH AUSTRALIA

COMPILED: R.A. Callen

LAKE FROME AREA - CAINOZOIC
CORRELATION CHART

SCALE: —

DRN: A.R.

CKD.

DATE:

PLAN NUMBER:

512169

TABLE 4

ROCK UNIT	MAP SYMBOL	LITHOLOGY	SEDIMENTARY STRUCTURES	CRITERIA FOR CORRELATION	FOSSILS	AGE	GEOMETRY	GEOMORPHIC EXPRESSION	SOIL DEVELOPMENT
STREAM DEPOSITS	Qra	Pebbles, sand, minor clay of modern stream bed. Silt and clay of low angle fans along western shore of Lake Frome, lower part may include fans of Coonabine Formation age.	Fine to medium scale cross-bedding, ripple marks, horizontal lamination, imbricated pebbles, etc. in streams.			Recent	Linear (stream channels) - (fans) (0-4m)	Bedload of streams, low-angle fans along west shore of Lake Frome.	
LAKE DEPOSITS	Qrl	Sand, silt, clay, silt. Black to green and brown	Unbedded or laminated, soft shows pressure ridges and blister structure.	Essentially continuous with Qrb clays.		Sub-recent	Thin spread (A few cms)	Windblown sand in lake bed, salt crust, silts and clays of mound springs.	?
UNWASHED SANDS	Qrs	Fine to medium sands. Bright red brown.	Aeolian cross-bedding - very large scale.	Unconsolidated, red-brown colour. Youngest unit in landscape.	Rare - insect carapaces, seeds, Charales oogonia.	Sub-recent	Thin linear (0-6m)	Modern windrift dunes of Strzelecki Desert.	?
SANDS	Qrb	Fine brown quartz sand. Blue green, brown and black structureless semi-liquid clay.	Sub horizontal bedding in sands.		Ostracods Charophytes (some reworked)	RECENT	Benched linear (0-4m)	Forms low benches around islands in Lake Frome and along eastern shore. False spits and bars of west shore may be equivalent.	?
DISCONFORMITY									
CALCAREOUS PALUDOSOL	Qec1	Cylindroids and patches of soft white carbonate, sometimes oriented in horizontal and vertical tubules or cylindroids. Usually poorly developed.		Generally less well developed than Qec2		PLEISTOCENE - RECENT? From Radiocarbon Dating	Benched linear	Caps benches	Probably the Bca horizon of a fossil soil.
COONABINE FORMATION	Qp1	Sand to silty or clayey sand, poorly to moderately sorted, usually fine-grained, with subangular to subrounded grains. 2.5-10YR5-8/4-6	Weak horizontal lamination, lenses, large scale cross-bedding (rarely well-defined)	Weakly developed bedding, moderate sorting, light brown colour, absence of gytticrete and large carbonate nodules. Locally abundant land snails. Typical soil horizon at top.	Snail shells, emu shell, aboriginal artifacts, Charales oogonia (rare), Rare vertebrate burrows.	PLEISTOCENE - RECENT?	Thin blanket (0-2m)	Flat low-lying plains, dominates landscape of planes and seif dunes. Crops out in creeks as vertical faces with columnar jointing.	Weakly developed carbonate cylindroids, weathering as hard granules. Discontinuous. Rectangular blocky structure forming columns of sediment.
	Qp2	Gypsum sand, coarse to very coarse, imbedded clay pebble layers. Lime biscuits and shell in basal layers.	Large scale cross bed sets. Complex fine scale cross-lamination.	Gypsum sand, moderate angle cross-stratification.	Charales oogonia Rare gastropods		Wedge shaped (0-5m)	Cliffs along southeastern shore of Lake Frome and islands in the lake.	Powdery gypsum crust with columnar jointing, which may persist into EURINILLA FORMATION. Qec1
DISCONFORMITY									
CALCAREOUS PALUDOSOL	Qec2	Soft white carbonate cylindroids tubules blotches and lumps in several layers, weather out as hard granules or crusts.				PLEISTOCENE - RECENT? From Radiocarbon Dating	Soil horizon		Probably the Bca horizon of a fossil soil.
UNWASHED GRAVEL	Qp3	Cobble and pebble deposits.	Imbricated flat pebbles in some localities.	Level 00-15m same as Qp2		Stalactites Dolomite Dolomite Dolomite	Linear Stalactites (0.5-2m)	Ridges parallel to western shore of Lake Frome	Weak carbonate cementation Qec1
GYPHRETE	Qcs	Gypsum crusts, of rosettes of disc-shaped crystals. Developed in various layers according to porosity. Includes calcareous cement in basal Eurinilla formation.				Associated with lower part of EURINILLA FORMATION	Discontinuous masses	Often forms massive cap, causing underlying strata to form in cliffs.	Groundwater origin - part same origin as younger Qca unit, with which it is associated in same horizon.
EURINILLA FORMATION	Qp4	Clay-sand to sand, with interbeds of silt and clay in upper part, lower part sand, often with pebbles. Very poorly sorted, grains may be moderately rounded, but rough and pitted, or rounded, and frosted. Upper part 2.5-5YR5-9/4-7 Sandy lower part 7-10YR5-7.5/2-6 Clays 5Y7/1-2.	None or crude horizontal stratification well-developed horizontal lamination (2-10 cm). Sometimes developed into large scale low angle sets. Basal sands rough or planar cross-bedded on medium scale. Rarely lower part horizontal wavy laminae with load casts.	Poor sorting, fine upper part, coarse lower part. Upper part lacks bedding or more rarely has well-defined sub-horizontal cross-stratification, lower part cross-bedded. Upper part reddish-brown, lower yellowish brown. Typical soil horizons, ground water calcareate and gytticrete (in basal part).	Insect burrows and nests. Vertebrates basal channels, including Diprotodon, Procoptodon sp. Basal layer of gastropods and lamellibranchs (Coxiella, etc) near edge of Lake Frome. Numerous charales oogonia, especially in coarse sand.	UPPER PLISTOCENE (Oneopus)	Thin tabular widespread sand sheets, thin hollows (0-40m)	Flat low-lying plains, cropping out as low slopes around depression and watercourses. Lower part often cemented by gytticrete or calcareate, producing bench or massive cap to cliffs. Frequently basal channels approximate present clay drainage.	Several well developed horizons of carbonate and gypsum cylindroids may be present at top, with large nodules in some areas. Frequently weather out as granules or sheets on surface. Lower sand cemented with crystalline white calcareate, weakly to strongly. Also with massive or layered gypsum rosettes. Mottled in shades of red-brown and white. Qec2, Qcs, Qca?
	Qp5	Pale greenish fine sands with wavy green clay laminae and coarse sand lenses. Thin basal lenses of bright orange brown sand. White sands at base of Eurinilla Formation. Covered by chara oogonia and shell fossils.	Small to medium scale cross-bedding. Clasts of clay.	Foraminifera, fish bones, greenish sand.	Charales stem moulds, oogonia, fish vertebrae and spines, Coxella, Foraminifera	EARLY PLISTOCENE or EARLY PLISTOCENE on boundary of EURINILLA FORMATION	Thin discontinuous sand to other units	Forms white benches near base of cliff profiles.	Affected by carbonate mottles patches in northern areas. Qec2, Qec1
POSSIBLE SAND (see EURINILLA FORMATION, unit Qp4)	Qca	White to pale brown secondary carbonate cement developed in coarse grained beds, also greenish carbonate with large cylindroids and tubules of carbonate. Black manganese staining in sands of topographic lows.		May include two calcareate Black stain may be related to older bed, but younger calcareate, younger calcareate equivalent to gytticrete.		One-sided with EURINILLA FORMATION on boundary of EURINILLA FORMATION	Thin discontinuous sand to other units	Solid outcrop, causes cemented units to stand out as prominent benches.	Groundwater calcareate.
UNDIFFERENTIATED CONGLOMERATE	Qp6	Undifferentiated conglomerate, grading to coarse sand. Intertongues with greenish brown silty clay in some areas.	Planar cross-bedding, trough cross-bedding, crude lenticular bedding.	constitutes several different units - conglomerates of Willawortina Formation, unnamed conglomerate Qp7, and conglomerate associated with Eurinilla Formation (Tel-ford Gravel equivalent)		Associated with EURINILLA FORMATION on boundary of EURINILLA FORMATION	Discontinuous (1-5m)	Solid outcrop, capping cliffs, or as ridges projecting above Lake Frome Plains.	Massive hard or soft, white, calcareate, cements coarser layers.
DISCONFORMITY									
MULLIVERA BEDS (see EURINILLA FORMATION, unit Qp4)	Qm	Pale green laminated clay, conchoidal fracture with thin Charales limestone (5Y9/1) laminae at top, interbedded with fine well-sorted sand (2Y7/4, 3Y7/2). Bedding absent in clay in Lake Frome. Limestone may grade to thin ripple-marked gypsum beds. Sand grains angular. Reddish to white fine sand with basal pebble beds and interbedded hard thin Charales limestone beds. Coarse loose white to greenish sands, sub-rounded grains. Algal limestone of mound springs. Clay: 10YR5/2, 5Y5/6, 5Y8-9/1-4, Fine sand: white to 2-3Y7/2-4. Channel sand: 5-10Y6-8/2, 5-10YR5-8/4-7.	Fine horizontal lamination in clays and limestone, asymmetric small scale ripples in gypsum, oriented Charales tubules, butyroidal blister structure on gypsum. Trough cross-bedded channel sands.	Charales tubule limestone, greenish calcareate-bearing clay, fine sand with Charales oogonia.	Charales stem moulds and oogonia. Spiral gastropods (Coxiella sp.), worm burrows (spiral), along bedding planes, with lined wood (pebbles, and tree trunks). Coniferous. Ostracods. Vertebrates in mound springs.	Uncertain: LATE MIOCENE - EARLY PLISTOCENE (from position relative to other units).	Thin low basin, clay on mound (5 m+)	In bed of Lake Frome and cropping out around edge, occupying old deep, etched stream channels, now represented by linear lakes associated with Willawortina - Eurinilla creek system. Mound spring limestone.	Massive white groundwater calcareate in sands, bright orange stain in basal coarse sands.
DISCONFORMITY									
UNWASHED CONGLOMERATE	Qp7	Massive calcareous conglomerate, calcareous sandstone. Basal cross-bedded light grey sand. Pinkish brown sand grains.	Planar cross-bedding, crude lenticular bedding.	Reworked carbonate nodules from Willawortina Formation. Interbedded yellowish brown clayey sand with greenish carbonate nodules.		Placed in late Pliocene - Pleistocene (from position relative to other units)	Linear (1-10m)	Solid outcrop capping cliff tops, particularly along Sicus - Ramore River system and southwestern shores of Lake Frome.	Massive white groundwater calcareate, often vuggy and crystalline - cements coarser beds. Qca

Note: Plan amended April 1976

FIG.8

DEPARTMENT OF MINES - SOUTH AUSTRALIA			
LAKE FROME AREA - CAINOZOIC QUATERNARY			
R.A. Cullen GEOLOGIST	Dm. R.A.C. Tot. A.F.	SCALE: —	74-772
Director of Mines	Ext.	DATE: SEPT '76	

Unit Qp₇ is possibly equivalent to the Telford Gravel of Firman (1963).

Upon these units is developed the calcrete horizon Qca, a white to pinkish massive carbonate which cements unit Qp₅ or sands of the Namba and Willawortina Formations. Iron and manganese oxide crusts, forming pipe-like structures and spherical bodies or sheets, are associated. The material is best developed east of Lake Frome.

Quaternary (Table 4, and large scale section D-D' on map).

The works of Williams (1971), Wasson (1979) and Callen & Tedford (1970) are useful for an understanding of the Quaternary discussed here.

Pleistocene

The Eurinilla Formation (Qp₄, Callen & Tedford 1976) has its type section at Lake Moko on *Coonarbine* (CURNAMONA) with supplementary sections at Lakes Millyera and Koorka. It is of late Pleistocene age, and contains basal channel sands with *Diprotodon* sp., *Procoptodon goliath*, and other species similar to those of the "Callabonna fauna". It is regarded as at least partly equivalent to the Tirari Formation (Stirton *et. al.* 1961).

The formation consists of fine to medium clayey poorly sorted sands, of orange brown colour, impregnated with gypsum at the base, and with interbedded green clays and silty clays. The upper portion often lacks sedimentary structures, or is horizontally laminated, whereas the lower portion is trough cross-bedded.

Excellent exposures are visible in extensive cliffs along the south side of Yandama Creek near its junction with Lake Callabonna, where the unit is overlain by the Coonarbine Formation. These are not shown on the map, having been discovered since its publication.

The unit has a strongly disconformable contact with the underlying units. East of Lake Frome it is generally cut into the Namba Formation or Millyera Formation and, west and south of the Lake, into the Willawortina Formation or unit Qp₇.

On FROME, undifferentiated gravels on the Lake Frome Plains which are included in unit Qp₆ occupy ridges in the landscape, though they are older than the surrounding sands (Qp₁, etc.). The reverse topography results from differential compaction and erosion. Most of the gravels may belong to the Qp₇ channel deposits, or unit Qp₄.

Old beach ridges (Qp₃), developed along the northwestern shore of Lake Frome, consist of gravel deposits and rare patches of material resembling unit Qp₅. The deposits are overlapped by the Coonarbine Formation sands. Recent evidence suggests the older (upper) ridges correlate with the unit Qp₅.

The next youngest unit is gypcrete (Qcs, map section D-D'), which forms massive to nodular crystalline sheets of wedge-shaped crystals or fibrous gypsum; it is present in the basal Eurinilla Formation or older units and is related to old land surfaces. Gypcrete frequently forms hard caps, permitting the development of finely patterned indented marginal cliffs along lake and watercourse edges which is readily distinguished from silcrete and usually also from calcrete.

Unit Qcc₂ a calcareous paleosol horizon, is typically developed on the Eurinilla Formation as widely developed horizons of soft white carbonate tubules and lumps, hard when weathered. Orange, brown and black ochreous mottling is associated. The unit appears as white patches on aerial photographs, difficult to distinguish from gypcrete.

Pleistocene - Holocene

The Coonarbine Formation (Qp₁, Callen & Tedford 1976)

sediments are moderately sorted sands, forming a thin widespread blanket east of Lake Frome. Cross bedding is present in some sections. Several mottled white carbonate paleosol horizons (Qcc_1) are present, weaker in development than the paleosol on the Eurinilla Formation (Qcc_2). The uppermost of these paleosols is frequently absent. A characteristic large scale rectangular jointing or pedality products a columnar weathering pattern. Locally, characteristic shells of the land snail *Thersites* (Buonaiuto op. cit.) are present. Radiocarbon dating of the paleosols demonstrates an age range from c. 30,000 to 7,000 yrs B.P., so that the unit extends back into the Pleistocene. Each paleosol is developed on an aeolian phase, the youngest of these being longitudinal dunes, which comprise the indurated cores of the Strzelecki Desert dunes. Loose sandy crests comprise the modern dunes (Qrs). Excellent outcrops, discovered since publication of the map, are present in the gully cut by overflow of Cootabarlow 3 bore.

Along the eastern shore of Lake Frome, and on its islands, are extensive gypsum/clay pellet lunettes (Qp_2 , see Bowler 1977). These discomformably overlie the Eurinilla Formation. Their extent is now known to be greater than recognised at the time of mapping - much of the material shown as Qp_4 around Lake Maljanappa is unit Qp_2 , and the lunettes along the edge of Lake Frome extend further north. Further east, each lake is flanked by a sandy leeside mound on its eastern margin, much of which is material of the same age as the gypsum lunettes. These have not been distinguished on FROME.

As mapped, the Coonarbine Formation includes some younger material of similar aspect, but without calcareous paleosol development. The paleosols may be absent in the upper unit of the Coonarbine Formation, and hence are indistinguishable from these deposits at the scale of mapping (lithology is otherwise

identical).

West of Lake Frome are extensive fluviatile sediments of very low angle fans and sheet deposits, which overlie the Eurinilla Formation. As they contain *Thersites* and have a similar appearance and weathering pattern to the Coonarbine Formation, they are included in this unit. They generally crop out as blocky 'breakaways' along creek edges. Calcareous soil is usually absent. Further work is needed to demonstrate their relationship to such units as the Mundi Mundi Unit of Wasson (1979).

On COPLEY and OLARY the units mapped as Pooraka Formation are essentially Coonarbine Formation. The Pooraka Formation as shown on adjacent maps constitutes the Coonarbine Formation and Eurinilla Formation combined.

The Coonarbine Formation has not been distinguished from unit Qrs east of Lake Frome (Thus the largest area shown as Qp₁ on the map is the alluvial material discussed above, and which recent evidence suggests should probably not be regarded as part of the Coonarbine Formation.

The Arrowie Formation (Coats 1973) on the Paralana High Plain constitutes brown sand with reworked cobbles, occasionally merging into thick cobble deposits derived from weathered Willawortina Formation, in which are land snails similar to those of the Coonarbine Formation. Some Eurinilla Formation is included, and also at least two phases of younger fans.

The Holocene dunefield (Qrs) of the Strzelecki Desert and part of the Lake Frome Plains consists of red brown loose sands and longitudinal dunes (Qrs, Simpson Sand of Firman 1970b in part). These have been largely reworked from the Coonarbine Formation dunes and Eurinilla Formation, and are partly active today. They have steep north faces, and are linked to transverse dunes along pan and lake edges. The unit dominates the eastern

half of FROME. As mapped, Qrs includes residuals of older Tertiary deposits, exposed in interdune areas, and the Coonarbine Formation. Eurinilla Formation is present along the water-courses.

Holocene beach deposits (Qrb) have been recognised, related to the present shoreline, and suggest a more active lake phase. False spits and bars are developed along the western shore, but are absent on the eastern side. They result from wind action on waters flooding into the lake from the Flinders Ranges. The deposits are probably related to the younger beach sands, and therefore are included in unit Qrb. Semi-liquid blue-green to brown and black clays of Holocene age (Qr1, Draper & Jensen 1976, Bowler 1977) form the uppermost playa sediments. They are distinguished from the Millyera Formation clays by absence of lamination. Low angle fans (Qra) occur along the western margin at creek outlets. A thin crust of salt is developed on the southern part of the lake (dashed line on map, Draper & Jensen 1976).

Since deposition of unit Qp₂, a gypsum karst structure has developed on the larger mounds along the eastern edge of Lake Frome. Modern gypsum/sand lunettes of small size are accumulating along the eastern edge of the present lake shore, and on some of the smaller lake shores.

STRUCTURE

Mt. Painter Inlier

The structural geology of the Mt. Painter Inlier has been described by Coats (Coats & Blissett 1971).

Rocks comprising the Mount Painter Complex form two discrete structural units, the Mount Babbage Block and Mount Painter Block. The Mount Painter Block was deformed prior to the Adelaidean into a series of easterly and northeasterly trending folds. Fold styles reflect competence of the rocks. Adelaidean

easterly trending domes and basins during an early Palaeozoic phase of tectonism. Faulting took place along pre-existing lines of weakness. The Mt. Adams Fault complex is a major structure, and is associated with breccia consisting mainly of Wywyana Formation. This and the marginal faults of the ranges were reactivated in late Cainozoic times. Some Mesozoic strata have been overthrust by Precambrian basement near Parabarana on CALLABONNA. *Concealed pre-Mesozoic rocks.*

The Benagerie Ridge coincides with a magnetic high, and is continuous with the well-defined structure on CURNAMONA (Jack 1930, Callen 1976a). In the southern part of the sheet, magnetic basement is interpreted at depths of 2-6km (Parker 1973, Tucker & Brown 1973). Tucker suggested that the Palaeozoic and Adelaidean sediments are not magnetic enough to account for the anomalies, which are better explained as intra-basement features. However, Bumbarlow 1 bore intersected red-beds and basalts of possible pre-Adelaidean age. The Mesozoic and lower Cainozoic strata (see cross sections) thin out over the Benagerie Ridge; Bumbarlow 1 and Mudguard 1 and private company uranium exploration drilling on *Eurinilla* (CURNAMONA), show the Middle Cambrian to be absent on top of this ridge.

A small coincident magnetic and gravity high along the eastern side of the Poontana structure in Lake Frome (Figs. 3, 4) is probably an intra-basement feature (pre-Adelaidean).

Seismic work by the South Australian Department of Mines and Crusader Oil N.L. indicates gentle folding in the Middle Cambrian rocks, with an overall westerly dip in the western third of the map area. Crusader's work indicates a narrow horst-like fault structure, encompassing probable Cambrian rocks, called the Poontana Structure (Coats *et. al.* 1969). This feature is not well-defined in the vicinity of Wooltana 1 bore. Topographic

features and drilling in the Cainozoic indicate it continues to the north where it forms the edge of the Paralana High Plain. This structure has been more accurately defined since publication of the Mt. Painter Special map, (Coats *et. al.* 1969) where it is shown as a fault or monocline to the west of its location on FROME. A similar narrow horst-like structure also marks the edge of the Balcanoona Highplain (COPLEY - see discussion and Fig. 3, Callen & Tedford 1977).

Frome Embayment and Tarkarooloo Basin

Configuration of the Mesozoic and Cainozoic basins differed. The Cretaceous basin has its axis beneath the eastern portion of Lake Frome (United Geophysical Corp. 1966), but during the Tertiary the locus of thickest accumulation moved westwards, closer to the edge of the present Flinders Ranges, resulting in the formation of the Poontana Sub-basin (unpublished Department of Mines and Energy plan 76/116 is available, showing contours for the base of the Cainozoic). This is illustrated on section A-A' on the map. In the Tarkarooloo Basin, fine grained sediments of the Namba Formation abut the bounding fault of the Flinders Ranges. At "Wertaloona" on COPLEY the Miocene sequence is represented by thin coarser sediments, and in the Munyallina Valley probable equivalents above the silicified Eyre Formation are coarse and feldspathic. This suggests the Flinders Ranges were present at the time, though of reduced height. Fan deposits must have ceased abruptly at the margins of the ranges. Intermontane basins were present, and have largely been stripped, but lacustrine sediments probably did not extend across the ranges. At the time, the ranges were probably thickly mantled with soil.

The Mesozoic and Cainozoic strata are virtually flat lying, though a westerly dip of $1-2^{\circ}$ west and south of L. Frome is inferred from drilling results; this is consistent with the dip

on the Mesozoic-Palaeozoic unconformity calculated by Horwitz (1958).

The Mesozoic and Cainozoic strata have a fault contact with the Adelaide "Geosyncline" sediments and Mount Painter Inlier igneous and metamorphic rocks lying to the west. The bounding fault system constitutes high angle reverse and possibly thrust faults, and coincides closely with the abrupt eastern margin of the Flinders Ranges.

Other structures

On the New South Wales border, steeply dipping quartz-veined ?Adelaidean (or ?Palaeozoic) rocks are exposed, implying that a basement structure affects the Cainozoic sediments.

Another structure is suggested by the alignment of mound springs along the eastern side of Lake Frome. Some of these may derive water from the Great Australian Basin (Draper & Jensen 1976). Sections AA', BB' illustrate a thickening and deepening of the Tertiary sequence west of this line.

The possible fault indicated at 140°33', crossing the southern border of the FROME map onto CURNAMONA, is interpreted from a ridge of silcreted Tertiary sediments, which has probably been upfaulted. The structure also corresponds approximately to the edge of a major magnetic feature on CURNAMONA separating a magnetic pattern typical of Willyama Inlier crystalline basement to the west from a linear northerly oriented magnetic low to the east. The latter is possibly the effect of gently folded Cambrian rocks. Extended north, the trend of the magnetic feature and fault passes between Mudguard 1 and Yalkalpo 1 bores. A thick sequence of Cambrian rocks lies to the east, and shallow Precambrian rocks to the west.

Draper & Jensen (1976) have suggested Lake Frome, like Lake Eyre (Johns & Ludbrook 1963, Dulhunty 1978), is tilted

to the south.

SEQUENCE OF EVENTS

1. In the Mt. Painter-Mt. Adams area, deposition of a thick sequence (6 100 m on COPLEY) of sediments began at least 2000 MY ago. This has since been metamorphosed to produce the Radium Creek Metamorphics. Widespread occurrence of porphyritic acid volcanics suggests a stable basement block prior to deposition of the Adelaidean System.
2. The sequence was folded, faulted and uplifted.
3. After prolonged erosion, the sediments of the Adelaide 'Geosyncline' were deposited. The Wywyana Formation marble indicates a phase of carbonate deposition, later involved in diapiric intrusions during early Sturtian and later times. The Fitton Formation and Bolla Bollana Formation represent a possible glacial interval. The remainder of the Adelaidean sequence is not represented on FROME.

The major portion of the Lake Frome area probably persisted as a stable cratonic block during deposition of the Adelaidean, as there is evidence for a thin undeformed Adelaidean and Cambrian cover. Here, the Precambrian is represented by shales, red-beds and conglomerates with interbedded basaltic lavas. There were extensive rhyolitic porphyry flows.

Lower Cambrian marine carbonate sedimentation took place, at least flanking the southern part of the Benagerie Ridge. During Middle Cambrian times a sequence of redbeds accumulated, directly related to those of the central Flinders Ranges. These are essentially continental fluvial and deltaic deposits, with a marine transgression represented by the Wirrealpa Limestone.

4. During Middle Cambrian to Ordovician time the Adelaidean and Cambrian rocks of the mobile belt were folded and intruded by granites along the western margin of the sheet area as a result of the Delamarian Orogeny (Thomson, 1969). Further east the basement acted as a stable cratonic nucleus (the Curnamona Cratonic Nucleus), where Delamarian folding is effectively absent.

5. The next depositional event was in Early Jurassic times, when epeirogenic movements initiated formation of the Great Australian Basin of which the Frome Embayment is a part. The Embayment comprises a lobe of sediments open to the main basin to the north. Crystalline basement areas were exposed in the Mount Painter area and along the Benagerie Ridge (Callen 1975).

Wopfner (1969) regarded the Parabarana Sandstone (on CALLABONNA, see Ludbrook, 1965) as representing the beginning of the Lower Cretaceous marine transgression following non marine Jurassic conditions. The Cadna-Owie Formation is thought to represent a nearshore facies, which intertongues with the lower part of the Marree Formation at the basin margins.

6. During the Late Cretaceous there was a major regression.

7. In the Early Palaeocene, renewed uplift initiated erosion of the regressive sediments of the Late Cretaceous. A new phase of non-marine sedimentation took place. A widespread sand blanket, the Eyre Formation, was laid down by braided and, subsequently meandering streams, over much of the Great Australian Basin. The basin had an outlet to the sea, and constant reworking and high rainfall caused the abundant fines eroded from the Cretaceous sediments to be flushed out. High rainfall and humid conditions are indicated by the floral content. Forests grew in the vicinity. Polished pebbles of fossil wood, chert and porphyry

originated from the Mesozoic strata.

The fine, brown, carbonaceous silts of the Eocene represent swampy or lacustrine conditions with meandering streams, indicating decrease in stream gradient as adjacent upland was eroded.

8. Widespread stable conditions then prevailed during the Oligocene, which was essentially a period of nondeposition. Silcrete developed sporadically along the margins of the Flinders and Barrier Ranges as a result of peculiarities in groundwater composition and climate, but is rare on FROME, which is in the centre of the Tarkarooloo basin.

9. In Medial Miocene times the Poontana trough began forming west of Lake Frome, and was the site of deposition of very finely laminated fissile carbonaceous clays of the lowermost Namba Formation, in a probable freshwater lake environment. The vertebrates and microflora indicate permanent water and high rainfall. Warm conditions are suggested from foraminiferal evidence in equivalent marine strata of the Murray Basin. High in the Miocene sequence, vertebrates are present. The taxonomic content implies a large body of permanent water with forested shores, in agreement with the high rainfall suggested from floral evidence.

Abundant smectite in the lower Namba Formation was formed under conditions of restricted leaching, though much was probably also inherited from the Marree Formation. An estuarine environment, or alkaline lake with internal drainage was present, and the palygorskite-dolomite sequence developed during times of seasonal aridity. Intermittent exposure of the lake sediments to weathering is recorded by incipient soil formation.

10. During deposition of the upper Namba Formation (unit Tmb₁), with its illiterich claymineral suite, a marked change in conditions occurred. This resulted from uplift of the Flinders Ranges in Upper Tertiary to Pliocene times, and imposition of a more seasonal, possibly cooler climate.

Such uplift was first suggested by Howchin (1913). Binks (1972) records a similar history for the Ediacara area (see also Callen & Tedford 1976 for discussion). The lakes were swamped with detritus, and extensive alluvial fans of the Willawortina Formation formed. Meandering streams dominated the lower part of the fan environment whilst mud flow and streamflood deposits built up near the ranges. These sediments were accommodated by the Poontana Subbasin. There is evidence for uplift, and weathering of high areas upon which alunite developed, before the Willawortina Formation was deposited.

11. In Late Tertiary times, after deposition of the Namba Formation, there was a period of extensive ferruginisation and silicification (CzSi).

12. This was followed by vigorous uplift of the Flinders Ranges, and to lesser degree, the Barrier Ranges and Olary Ranges.

13. Deposition of the Millyera Formation lacustrine, fluviatile and beach deposits followed. In the Pasmore River region, extensive conglomerates (Qp₇) probably represent the ancient river system feeding this lake. The channels roughly coincided with present day rivers. Rainfall must have been higher than at present as the lake was much larger and extended

towards the New South Wales border. Channel deposits were extensive. Evaporation was high toward the end of this phase.

14. Calcrete and mottled ferruginous paleosol developed.

15. The Eurinilla Formation records an essentially fluviatile environment, dominated by meandering streams, closely coincident with present drainage. Lake Frome decreased to something approaching its present size. Large marsupials lived in the vicinity. The higher beach deposits west of Lake Frome were probably formed about this time (?c.50,000 years B.P.)

16. A calcareous paleosol then formed over a wide area.

17. In Late Pleistocene to Early Holocene times, a new phase of deposition set in, represented by the Coonarbine Formation, which is essentially aeolian. An extensive system of seif dunes and gypsum lunettes developed, in at least three phases. During the major Quaternary aridity (c.25 000 to 18 000 B.P.), lunettes formed in and adjacent to the eastern side of Lake Frome, and smaller lakes to the east. Dry cold westerly winds prevailed. Carbonate paleosol horizons are intercolated between the aeolian phases. Fan deposition occurred west of Lake Frome, but it is not known whether this coincided with dune development; it is most likely to have alternated with or postdated the major aeolian phases, perhaps coinciding with interglacials. Younger beach deposits are now known from the western shore of Lake Frome.

During Late Pleistocene times the Flinders Ranges and flanking High Plains continued to be uplifted, whilst the Olary

and Barrier Ranges were comparatively stable. Lacustrine sediments continued to be deposited in Lake Frome.

ECONOMIC GEOLOGY

Metallic Minerals

The mineral resources of the Mount Painter Province have been discussed by Blissett (Coats & Blissett 1971). Potential for copper and uranium mineralisation is considered to exist along the Mount Adams fault system. The Gunsight Prospect lies on this structure at the northern margin of the sheet.

Reserves of 15,900 tonnes of uranium oxide have been proved at the Beverley by Western Nuclear Aust. Ltd. in joint venture with the Transoil-Petromin-Oilmin group. The uranium occurs in medium grained fluviatile sands in the uppermost part of Member 2 of the Namba Formation. The Poontana structure controls sedimentation, and has channelled the uraniferous groundwaters which led to deposition of the uranium. On CURNAMONA, small prospects occur in channels of the Eyre Formation but as yet no uranium has been found in this unit on FROME. Economic aspects of this geology are discussed by Callen (1974a, 1976a), Brunt (1978) and Haynes (1976) and extractive leaching by Lackey (1974).

Heavy minerals with a high percentage of titanium occur in the Eurinilla Formation on the western shore of Lake Frome. Other deposits containing 70% zircon and 2.8% rutile in the heavy fraction were found in surface deposits at Lake Cootabarlow (Stillwell 1949). The sands of the Eyre Formation, particularly on CURNAMONA, contain up to 5% rutile, zircon and other heavy minerals.

Rare manganese "wad" that occurs as massive layers in the Miocene sands is related to groundwater movements in the Plio-Pleistocene. Outcrop of this material has been prospected along Tielta Creek.

Non Metallic Minerals & Fossil Fuels

Swelling clays of the palygorskite group (a variety of Fuller's Earth) and degraded ?montmorillonite type occur in association with dolomite, in outcrops around many of the small lakes. Demand for palygorskite for use as industrial absorbants and catalysts is such that exploration has begun on CURNAMONA immediately southeast of Lake Frome (Barnes & Olliver 1978). Thick white kaolinite deposits are associated with sands of the Eyre Formation particularly adjacent to granitic source areas, and have been investigated in the Paralana High Plains on COPLEY.

Extensive impure gypsum is present, usually mixed with clay and sand in aeolian lunettes, and also as thin sheets in Lake Frome. The Namba Formation has some similarities with the Green River Formation, and hence has potential for evaporites, exploration by Comalco Exp. Corp. is proceeding.

Alunite, occurs over a wide area in the upper part of the Namba Formation, but is often impure, being mixed with clays. Outcrop occurs at Lakes Bumbarlow and Starvation, and has been recorded in bores elsewhere.

The Eyre Formation contains large fragments of coalified wood (e.g. Yalkalpo 1 bore) and methane gas has been detected from bores in the eastern side of the lake, but no significant fossil fuels are known. From her studies of the Wirrealpa and Aroona Creek Limestone, Youngs (1978c) has shown the Arrowie Basin has potential for hydrocarbons north and northeast of Lake Frome (Youngs 1978a,b).

Hydrogeology

Several workers have studied the hydrogeology, the most recent being Ker (1966). Underground water supplies are available over most of the area; artesian supplies are provided from

the basal Great Australian Basin sediments where these are present.

West of Lake Frome, water of variable quality is obtained from several sources including stream gravels of the Eurinilla and "Coonarbine Formation", unconsolidated sands of the Willawortina Formation channels in the northern part of the area, and blanket Tertiary sands which provide moderate stock quality sub-artesian water.

East of Lake Frome, the Tertiary aquifers provide good supplies (Brown 1950, who misidentified the Eyre Formation aquifer as Cretaceous). The major aquifer, the Cadna-Owie Formation or Algebuckina Sandstone (Interstate Conference on Artesian Water 1929), has not been used west of Lake Frome. Springs situated along the western shore of Lake Frome and the mound springs within the lake are too saline for stock.

Flow from the eastern artesian bores (Boucaut & Beal, 1977) is essentially uncontrolled and local graziers have noted a reduction of 25-50% during the past 40 years.



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OPEN FILE COMPANY REPORTS HELD IN THE ENVELOPE SYSTEM,
DEPARTMENT OF MINES AND ENERGY, SOUTH AUSTRALIA

Company	Date	Locality & map sheet, minerals	LEASE	ENV.	Rock strat. units
Afmeco Pty. Ltd.	1972	Lake Frome, SH54-10, 6837-3, U	SML634	2016	C
Afmeco " "	1973	Lake Frome, SH54-10, 6837, 6937, U	EL19	2176	C
Afmeco " "	1973	Beverley Area, SH54-10, 6837-4, U	EL64	2535	C
Oilmin NL., Transoil NL., Petro- min NL., & Afmeco Pty. Ltd.	1977	Poontana Sub-basin, SH54-10, SH54-9, 6937-3, 4, 6737-3,4, 6837-2, U.	EL176	2535	T
Anaconda Australia Inc.	1966	Paralana Springs, SH54-9, SH54-10, SH54-6, SH54-5, 6737-1,4, 6837-4, 6838-2, 6738-2,3, Base metals.	SML112	633	B (stream sed.)
Central Pacific Minerals NL., & Magellan Petroleum (N.Y.) Pty. Ltd.	1971	Lake Frome, SH54-10, 6837, 6937, U	SML462	1746	C
Central Pacific Minerals NL., & Magellan Petroleum (N.Y.) Pty. Ltd.	1971	Lake Frome SH54-10, SH54-9, 6937-3,4, 6737-2. U	SML240	1408	C,M
Chevron Exploration Corporation	1972	Quinyambie Prospect, SH54-10, 7036 7037-2,3, U	SML663	2100	T
Chevron Exploration "	1973	Quinyambie Prospect, Sh54-10, 7037-2,3 7036, U	EL40	2257	T
Delhi International Oil Corp., Santos Ltd. & Vamgas N.L.	1970	Frome Downs SH54-5, SH54-6, SH54-9, SH54-10, SH54-14, SH54-14, 6738, 6838, 6938, 7038, 6737, 6735. Petroleum	PEL5/6	1124	M.C. Cam. (Wirrealpa L/S)
Delhi International Oil Corp., Santos Ltd. & Vamgas NL.	1971	Frome Downs, SH54-10, SH54-13, SH54-9 SH54-5, SH54-14, 6735, 6736, 6737, 6738, 6835, 6935. Petroleum.	PEL5/6	1275	M Cam. (Wirrealpa L/S)

Company	Date	Locality & map sheet, minerals	LEASE	ENV.	Rock strat. units
Delhi International Oil Corp., Santos Ltd., Vamgas N.L.	1971	Frome Downs SH54-10, SH54-13, SH54-14, Sh54-6, 6735, 6736, 6835 6935, 6838, 7035. Petroleum	PEL5/6	1566	M,C. Cam. (Wirrealpa L/S)
Delhi Australian Petroleum Ltd. & Santos Ltd.	1963	Lake Frome, SH54-9, SH54-10, SH54-13 SH54-14, 6736-2,3, 6836, 6735, 6835-34, Oil, Gas.	OEL20/21	287	K-Cam.
Delhi Australian Petroleum Ltd. Santos Ltd.	1963	Lake Frome, SH54-13, SH54-9, SH54-14 SH54-10, SH54-6, 6635, 6735, 6736, 6737-1, 6835, 6836. Petroleum.	OEL20/21	305	
Delhi Australian Petroleum Ltd. & Santos Ltd.	1963	SH54-13, SH54-9, SH54-10, SH54-6, SH54-14, 6735, 6736, 6837, 6937, 6836, 6737-2. Petroleum.	OEL20/21	306	
Delhi Australian Petroleum Ltd. & Santos Ltd.	1963	Lake Frome, SH54-13, SH54-9, SH54-5 SH54-14, SH54-10, 6735, 6736, 6737 6738, 6835, 6837, Petroleum.	OEL20/21	335	
Delhi Australian Petroleum Ltd. & Santos Ltd.	1965	Strzelecki Desert. SH54-10, SH54-13 SH54-14, SH54-2, SH54-10, SH54-9, SH54-1. Petroleum.	OEL20/21	569	
Delhi Australian Petroleum Ltd. & Santos Ltd.	1965	Frome Embayment, SH54-10, SH54-13 SH54-14, SH54-1, SH54-2, SH54-5, SH54-5, SH54-10, Petroleum.	OEL20/21	510	
Delhi Australian Petroleum Ltd. & Santos Ltd.	1966	Eromanga, Cooper Basin, SH54-13 SH54-14, SH54-9, SH54-10, SH54-6, SH54-5, SH54-1, SH54-2, SH54-15, SH54-10, 6735, 6835, 6736, 6836, 6739, 6740, Oil, gas.	OEL20/21	556	P
Kerr McGee Australia Ltd.	1969	Mt. Painter area, SH54-10, SH54-6, 6837-1,4 6838-2,4, U.	SML218	986	C

Company	Date	Locality & map sheet, minerals	SML	ENV.	Rock strat. units
Mines Administration Pty. Ltd.	1970	Lake Callabona. SH54-6, SH54-10, 6837, U	SML431	1440	T
Mines Administration " "	1971	Lake Frome, SH54-10, SH54-14, 6835, 6836, U.	SML531	1622	C,M.
Mines Administration " "	1972	Lake Callabonna, SH54-10, SH54-6, 6837-1, 6838-1,2, 6937-4, 6938-3, U.	SML596	1800	T
North Flinders Mines N.L.	1970, 1971	Coup Prospect, Arkaroola, Wooltana SH54-10, 6837-2,3, Cu, Zn	SML294	1226	B(Blue Mine Conglomerate, Wywyana F., Ulupa Siltstone, Wooltana Volcanics, Opaninda F., Humanity Seat F., Wortup Quartzite, Skillogalee Dolomite)
North Flinders Mines N.L.	1969	Willigan Hill, Parabarana, Mt. Adams, Mt. Fitton, SH54-9, SH54-10, SH54-5 SH54-6, 6737-1, 6837-4, 6738-2, 6838-3, Cu, U	SML291	1224	B
North Flinders Mines N.L.	1969, 1970, 1971	O'Donoghues Castle, Wooltana, Mcleashes Mine, SH54-10, SH54-6, 6837-2,3. 6838-3 Cu, Pb, Zn.	SML294	1226	B (Woodnamoka phyllite, Callanna Beds) (stream sediment & soil samples) Adelaidean (Burra Group Blue Mine Conglomerate, Skillogalee Dolomite).
North Flinders Mines N.L.	1971	Pinnacles, Mine, Yudnamutana Mine,	SML558	1369	Adelaidean, B.
North Flinders Mines N.L.	1970, 1971	Mt. Adams SH54-10, SH54-6, 6837-4, 6838-3. Parabarana, U, Cu, Pb, Zn.	SML296	1220	B (soil sampling)
North Flinders Mines N.L.	1971	Freeling Heights, SH54-9, SH54-10, 6737-1, 6837-4, base metals	SML558	1639	B (stream sed.)
North Flinders Mines N.L.	1971	Moolawatana, SH54-5, SH54-6, SH54-9, SH54-10, 6738, 6838, 6737, 6837, Cu, Zn, Pb.	SML444	1442	B (stream sed.)

Company	Date	Locality & map sheet, minerals	SML/EL LEASE	ENV.	Rock strat. units
North Flinders Mines Ltd.	1971 1972	Pinnacles Mine, Yudnamutana Mine, SH54-5, SH54-6, SH54-9, SH54-10 6738-1,2, 6838-3,4, 6737-1, 6837-4 Metalliferous Minerals, Cu, U.	SML558	1369	Adelaidean, B. (stream sediment and soil samples) Fitto F., Carpentarian, Terrapinna Granite.
North Flinders Mines Ltd.	1972	Mt. Painter Province, Gunsight, Carthew, Brindana Mine, Pinnacles Mine, Mt. Neil, SH54-10, SH54-6, 6837-4, 6838-3. Cu, U.	SML558	1369	B (Rock chip sampling)
North Flinders Mines Ltd.	1972, 1973, 1974	Gunsight, Mt. Neil, Mt. Babbage, Mt. Painter Province, Carthew, Mt. Adams, Pinnacles, SH54-6, SH54-9. SH54-10, 6737-1,4, 6838-3, 6837-4. Base metals and U, Cu, Zn, Pb, Fe.	SML704, 588	2034	B (stream sed., etc)
Billy Springs Pty. Ltd.	1968	Mt. Adams, SH54-10, 6837-4, base metals	SML194	1038	B
Oilmin N.L., Transoil N.L. & Petromin N.L.	1979	Lake Yalkalpo, SH54-10, 6936-2. U.	EL386	3315	C
Delhi Australian Petroleum Ltd. & Santos Ltd.	1959	Frome Embayment, SH54-6m SH54-10 Oil, Gas.	OEL20/21	69	M
Delhi Australian Petroleum Ltd. & Santos Ltd.	1960	Frome Embayment SH54-10, 7037 Oil, Gas.	OEL20/21	65	J
Sedimentary Uranium N.L.	1972	Lake Frome, SH54-10, 6936, U.	SML630	1827	T?
Southern Ventures Pty. Ltd.	1974	Lake Carnanto, SH54-10, 7036-2,3. U.	EL105	2377	C
Southern Ventures " "	1974	Lake Yalkalpo SH54-10, SH54-14, U.	EL90	2467	C,K
Southern Ventures " "	1974	Lake Carnanto, SH54-10, U.	EL105	2468	C,K
Southern Ventures " "	1974	Lake Yalkalpo, SH54-14, SH54-10, 7035, 7036-3,4. U.	EL90	2327	
Southern Ventures " "	1976	Lake Carnanto, SH54-10, SH54-14, 7036, U.	EL178	2584	T

Company	Date	Locality & map sheet, minerals	EL LEASE	ENV.	Rock strat. units
Southern Ventures Pty. Ltd.	1977	Quinyambie, SH54-10, SH54-14, 7036-2,3, EL334 7035-1,4. U.		3058	T
Tricentrol Australia Ltd.	1974	Lake Elder Area, SH54-10, 6936-2,3, U.	EL34	2392	T,M.
Tricentrol " "	1974	Lake Elder, SH54-10, 6936, U.	EL34	2211	T.
Tricentrol " "	1975	Lake Coonarbine, SH54-10, SH54-14, 6936-2,3, 6935-1,4, U.	EL127, (69,34)	2432	T.
Union Corporation (Australia) Pty. Ltd.	1973	Frome Embayment, SH54-10, 6937, 7037, U.	EL5/6	2338	T,K.
Union Corporation (Australia) Pty. Ltd.	1973	Quinyambie, SH54-10, 6937, U.	EL6	2146/2338	C,K.
Union Corporation (Australia) Pty. Ltd.	1973	Lake Pundalpa SH54-10, 6937-2,3, 7037-2,3. U.	EL5	2145/2338	T
Western Mining Corp. Ltd.	1978	Wittakilla SH54-6, SH54-10, 6838 6938, 6837, 6937, base metals	EL364	3153	T,K. B not interested

Note: Symbols used under column Rock Strat. Units.

B = Crystalline basement
 C = Cainozoic
 K = Cretaceous
 M = Mesozoic
 Cam = Cambrian
 J = Jurassic
 T = Tertiary
 P = Permian